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AN UNPREDICTED ANCIENT COLONIZATION OF THE WEST INDIES BY NORTH AMERICAN RODENTS: DENTAL EVIDENCE OF A GEOMORPH FROM THE EARLY OLIGOCENE OF PUERTO RICO

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Abstract: West Indies are a natural laboratory for the study of biogeography and evolution, especially for mammals. The modalities of their arrival in the Caribbean islands are highly controversial issues, and palaeontological evidence remains particularly elusive. Lower Oligocene deposits of Puerto Rico (c. 29.5 Ma) have recently yielded the oldest archipelago rodents, with teeth of two chinchilloid caviomorphs of South American origin. Here we report the discovery of additional dental remains from the same level, testifying to the occurrence of another rodent group. Their morphology is singular but otherwise characteristic of North/Central American geomyin geomorphs (including extant kangaroo rats, pocket mice, and pocket gophers), and leads us to propose here a new taxon: *Caribeomys merzeraudi* gen. et sp. nov. The evidence of geomorphs in the West Indies challenges the view that all extinct and extant Antillean rodents derived from South American sources. The morphological singularity of this Puerto Rican geomorph also suggests an early split of this insular lineage from coeval mainland counterparts, and its dispersal towards the West Indies sometime around the Eocene–Oligocene transition. This unexpected rodent adds to the few cases of Antillean terrestrial vertebrates of North American origins, such as solenodonotan eulipotyphlans. In the absence of subaerial land connections between both landmasses, an overwater dispersal may explain mid-Palaeogene colonization of the West Indies by geomorph rodents.

Key words: Caribbean islands, Rodentia, Geomorpha, Palaeogene, palaeobiogeography.
RODENTS are one of the most speciose land mammal components of the Caribbean islands (e.g. Wilson & Reeder 2005). Regardless of the many species introduced by humans (e.g. Cooper 2008; MacPhee 2009), hutias (Capromyinae, Echimyidae, Octodontoidea, Caviomorpha) and rice-rats (Oryzomyini, Sigmodontinae, Cricetidae, Muroidea) were once the most diverse West Indian rodent groups, although representing only part of the past rodent diversity on these islands. All rice-rats and several species of hutias have recently become extinct (e.g. MacPhee 2009; Brace et al. 2015; Cooke et al. 2017; Turvey et al. 2017; Upham & Borroto-Páez 2017), but so have other endemic caviomorph hystricognaths, such as the heteropsomyine echimyids and the iconic “giant hutias” of chinchilloid affinities (see Marivaux et al. 2020). Their phylogenetic origins, historical biogeography and the causes of their extirpation, have long been debated issues. Comparative anatomy studies and phylogenies derived from available morphological and molecular data have shown that all extinct and extant rodent groups endemic to the Greater and Lesser Antilles, evolved from South American sources (e.g. MacPhee & Iturralde-Vinent 1995, 2005; Woods et al. 2001; MacPhee 2005; Fabre et al. 2014, 2017; Brace et al. 2015; Courcelle et al. 2019; Marivaux et al. 2020). Recent palaeontological evidence (Marivaux et al. 2020) and molecular divergence estimates (Fabre et al. 2014; Brace et al. 2015; Courcelle et al. 2019) have further revealed that the arrival of rodents on the West Indian archipelago likely occurred through multiple and time-staggered dispersal events (chinchilloids then subsequently echimyid octodontoids [+] caviid, but in southern Lesser Antilles only; MacPhee et al. 2000) among caviomorphs, and lastly oryzomyin muroids). Although the timing of these dispersals is now relatively better bracketed, the modalities (land connection versus overwater transport) and pathways of these dispersals remain hypothetical and disputed (e.g. Hedges et al. 1992; MacPhee & Iturralde-Vinent 1995, 2005; Hedges 2001, 2006; Dávalos 2004; Ali 2012). The palaeobiogeographical scenarios for land-mammals derive from limited palaeontological data and phylogenetic inferences (either based on morphology or on genes/proteins of living and recently extinct species), but also from partial and varying geological and tectonic models, especially regarding emerged landmasses (e.g. Pindell & Barrett 1990; MacPhee & Iturralde-Vinent 1995, 2005; Hedges 1996; Iturralde-Vinent & MacPhee 1999; Woods et al. 2001; Graham 2003; Roca et al. 2004; MacPhee 2005; Iturralde-Vinent 2006; Fabre et al. 2014; Brace et al. 2015, 2016; Courcelle et al. 2019; Delsuc et al. 2019; Presslee et al. 2019; Marivaux et al. 2020). The same is true of scenarios for the origins of many other terrestrial organisms on the Caribbean islands (e.g. Heinicke et al. 2007; Alonso et al. 2012; Matos-Maravi et al. 2014; Dziki et al. 2015; Chamberland et al. 2018; Čandek et al. 2019; Tong et al. 2019; Blackburn
et al. 2020; Crews & Esposito 2020). The current diversity of West Indian mammals is so eroded, and the fossil record so incomplete, that the picture of their past-diversity can only be sketchy.

Here we report the discovery of dental remains of a new rodent from the early Oligocene of Puerto Rico (c. 29.5 Ma; San Sebastian Formation). The fossils were collected in 2019-2020 in the residues of fine wet-screenings (1 mm ≤ grains < 2 mm) of sediments from the LACM Loc. 8060 fossil-bearing layer (Río Guatemala, San Sebastián; Fig. 1). In 2019, excavations of the same layer and coarse wet-screenings (> 2 mm) of the extracted sediments had already led to the discovery of dental remains of two chinchillloid caviomorph rodents (Marivaux et al. 2020), representing the earliest direct evidence of colonization of the Greater Antilles by South American rodents (Vélez-Juarbe et al. 2014). We describe here two isolated molariform teeth documenting a small rodent species, whose characteristics differ from those of any known muroids or South American caviomorphs, extinct and extant, but rather match those of some Palaeogene rodents recorded in North and Central America, known as Geomorpha (e.g. Flynn et al. 2008). This material is admittedly limited, but the morphology of these teeth leads us to propose here a new taxon: Caribeomys merzeraudi gen. et sp. nov. We analyse and compare this new rodent, and discuss the macroevolutionary and palaeobiogeographical implications of this ancient unexpected rodent colonization from North America to the Greater Antilles.

MATERIAL AND METHODS

Fossil recovery and extraction. The LACM Loc. 8060 fossil-bearing locality is a 15–20 cm thick layer of silty claystone (Vélez-Juarbe et al. 2007, 2014; Vélez-Juarbe & Domning, 2014; Marivaux et al. 2020). We excavated and collected c. 400 kg of sediments along this layer (Fig. 1). When fully dry, the gross sediments were treated by wet screening (two meshes of sieves; 2 mm and 1 mm) directly in the water of the Río Guatemala. The coarse residues of screening (≥ 2 mm) were carefully scrutinized by naked eyes in situ (on the riverbank), in order to collect the medium-sized fossil elements. The fine residues (1 mm ≤ grains < 2 mm) were generally screen/washed twice, dried, and observed under stereomicroscopes in our field laboratory, and afterwards. The fossil material consists primarily of isolated teeth, among which the two specimens of rodents reported here.
Fossil repository. The fossil specimens described in this paper are permanently stored in the palaeontological collections of the Natural History Museum of Los Angeles County (LACM), Los Angeles, USA.

Dental loci, nomenclature and measurements. The dental terminology employed here (Fig. 2) is updated and modified after Wahlert (1983), Korth (1994, 2007), Marivaux et al. (2004), Calede & Glusman (2017), and Calede & Rasmussen (2020). Teeth were measured (maximum mesiodistal length and maximum buccolingual width) with a microscope fitted with a calibrated reticle (Measuroscope Nikon 10).

High-resolution μCT-scan. We used X-ray microtomography (μCT-scan) to obtain three-dimensional digital models (3D renderings) of the fossils (Figs 2–4). The specimens were scanned with a resolution of 5 µm (LACM 162478) or 4.5 µm (LACM 162449), using a μCT-scanning station EasyTom 150 / Rx Solutions (Montpellier RIO Imaging, ISE-M, Montpellier, France). The crown and roots of each tooth were virtually delimited by manual segmentation under AVIZO 7.1 software (Visualization Sciences Group). The teeth were prepared within a “Label Field” module of AVIZO, using the segmentation threshold selection tool. The renderings of the 3D digital models of the teeth described in this paper (Marivaux et al. 2021) are available on the MorphoMuseuM website (http://morphomuseum.com), an online open access platform that aims at enhancing accessibility to 3D models.

Optical photography. Some of the tooth pictures (Figs 3A, 4A) are photographs of original specimens, whitened using magnesium smoke. These views are the result of the fusion of multifocus images obtained with an optical stereomicroscope (Leica M 205C) connected to a camera (Leica DFC 420C).

Institutional abbreviations. Amherst, Amherst Museum; F:AM, Frick American Mammals, the American Museum of Natural History; LACM, Natural History Museum of Los Angeles County, Los Angeles; MCZ, Museum of Comparative Zoology, Harvard; UMPE, Universidad del Mar, campus Puerto Escondido; UNSM, University of Nebraska, State Museum, Lincoln; SDSM, Museum of Geology, South Dakota School of Mines and Technology, Rapid City.
SYSTEMATIC PALAEONTOLOGY

We follow here the systematics of Geomorpha proposed by Flynn et al. (2008).

Class MAMMALIA Linnaeus, 1758
Order RODENTIA Bowdich, 1821
Suborder ISCHYROMYIFORMES Marivaux, Vianey-Liaud & Jaeger, 2004
Infraorder GEOMORPHA Thaler, 1966
Parvorder GEOMYINA Bonaparte, 1845
Family indeterminate
Genus CARIBEOMYS nov.

LSID. urn:lsid:zoobank.org:act:E55A5360-1CCB-42BD-A372-D858DA588600

Type species. Caribeomys merzeraudi sp. nov.

Etymology. The name derives from the Spanish “Caribe” (El Caribe) corresponding to the Caribbean region (Caribbean Sea and West Indies), with the Greek suffix μυς (mys), mouse.

Diagnosis. As that of the type and only species.

Geographic distribution and stratigraphic range. Puerto Rico (West Indies), late early Oligocene.

Caribeomys merzeraudi sp. nov.

Figures 3–4

**Etymology.** The species name pays tribute to our late colleague and friend, Gilles Merzeraud (1965–2020), who established the stratigraphical section of Rio Guatemala from which the fossils described here originate.

**Holotype.** LACM 162449, right lower molar (m1 or m2; Figs 2A, 4A–P).

**Hypodigm.** In addition to the holotype, LACM 162478, a right fourth lower deciduous premolar (dp4; Figs 2B, 3A–N).

**Horizon and locality.** Late early Oligocene (c. 29.5 Ma) San Sebastian Formation at locality LACM Loc. 8060, on the west bank of Rio Guatemala, San Sebastian, Puerto Rico (Vélez-Juarbe *et al.* 2007, 2014; Vélez-Juarbe & Domning, 2014; Ortega-Ariza *et al.* 2015; Marivaux *et al.* 2020).

**Diagnosis** (based on the two dental specimens available). Small geomyin species (c.100 g) with mesodont teeth displaying an unusually thick enamel layer. Differs from other Geomyina (i.e. Heliscomyidae, Florentiamyidae, Geomyoidea *incertae sedis*, Heteromyidae and Geomyidae) in having lower teeth with the buccal protostylid and hypostylid linked to the protoconid (or metalophid) and hypoconid (or hypolophid), respectively. Differs from Heliscomyidae (most species of *Heliscomys*) in having teeth much less brachydont and bunodont, without well-differentiated main cuspsids and buccal stylids. Differs from Florentiamyidae (e.g. *Kirkomys, Florentiamys*), some species of Heteromyidae (e.g. *Eochaeotodipus, Perognathus*) and all Geomyidae in having well-defined and buccolingually-extended anterocingulid and posterocingulid. Differs from some other low-crowned heteromyids (such as *Proheteromys*) and some stem geomyoids (e.g. *Proharrymys*) in having lingual and buccal cuspsids which, when merged (with wear), form transverse metalophid and hypolophid, not V-shaped lophids. Differs from most geomyids and high-crowned heteromyids in having teeth much more low-crowed and mesodont (not hypsodont or hypselodont).

Body-mass estimate of c.100 g is based on the LACM 162449 molar area, calculated from the least-squares regression equation of Legendre (1986).
**Description**

The two specimens are right lower teeth, LACM 162478 representing a pristine fourth deciduous premolar (dp4; Fig. 3A–N), and LACM 162449 a worn first or second molar (m1 or m2; Fig. 4A–P). The premolar is half the size of the molar, low-crowned, and its occlusal surface lacks any mark of abrasion and attrition wear, thereby indicating it is a dental germ that belonged to a new-born individual. In contrast, given its advanced stage of wear (abrasion related to mastication), the LACM 162449 molar belonged to an old adult individual. Both teeth display a noticeably thick enamel layer (Figs 3B–D, 4A–D), which has certainly contributed to their preservation (without distortion) during fossilization. The roots are lacking on both specimens, and crowns are preserved from the cervix. The basalmost part of the crown of LACM 162478, where roots should have started, displays small but well-marked mesial and distal outward folds (Fig. 3G, K), thereby indicating that two roots were present and anchored in the dentary with diverging directions on a mesiodistal axis. Such a root configuration strengthens support for the deciduous status of LACM 162478, as this orientation of the roots is usual for the unique lower deciduous premolar in rodents. The dentine is only partially preserved and highly fractured on the molar (Fig. 4I), whereas it has entirely vanished on the deciduous premolar, the crown of the latter consisting only of a hollow enamel cap (Fig. 3F).

*LACM 162478 dp4*. In occlusal view, this tooth is longer than it is wide (mesiodistal length: 1.51 mm, buccolingual width: 1.26 mm), having an overall rectangular shape but with rounded edges and corners (Fig. 3A, E, K, I, M). The trigonid region is only slightly narrower than the talonid region, making the occlusal outline of this tooth almost ovoid, with a mesiodistal long axis. Being pristine, this deciduous premolar reveals its low crown elevation (mesodonty; Fig. 3H, L), and its occlusal relief returns a somewhat singular structural pattern. Four main cuspsids can be identified, but only the mesialmost one, i.e. the metaconid, is well defined (Fig. 3A, E, G, I, M). The protoconid, hypoconid and entoconid appear more like discreet enamel swellings, subsumed within transverse cristids and accessory enamel wrinkles. The metaconid is protruding and mesiolingually canted, and is located slightly more buccal than the entoconid on a mesiodistal axis. The entoconid seems mesiodistally extended and buccolingually compressed due to the development of short anterior and posterior arms (Fig. 3A, E, G, K). The protoconid and hypoconid are mesiodistally aligned. The lingual
cuspids (metaconid and entoconid) are marginally positioned, whereas the buccal ones (protoconid and hypoconid) are internal, lingually offset from the buccal margin. The entoconid and hypoconid are buccolingually aligned and linked by a thin but trenchant hypolophid, which is transverse in its lingual part and slightly buccomesially oriented before reaching the hypoconid. The protoconid is distally positioned with respect to the metaconid. A very short metalophid, notched in its middle, links the distobuccal flank of the metaconid to the lingual aspect of the protoconid (Fig. 3A, E, G, M). The buccal margin is bounded by a moderately elevated buccal cingulid that is interrupted at its midpoint by a V-shaped notch corresponding to the buccal opening of the hypoflexid (Fig. 3G–H). This buccal cingulid bears well-defined and strong protostylid and hypostylid, which are located directly buccal to the protoconid and the hypoconid, respectively. An accessory stylid also occurs distal to the protostylid. The hypostylid and hypoconid are linked together by a very short and buccolingually oriented cristulid. A similar cristulid is also present between the protostylid and the protoconid. The buccal cingulid is prolonged mesiolingually by a low, short but well-marked and arcuate anterocingulid, and distally by a low but well-defined, long and slightly circular posterocingulid. The anterocingulid connects the protostylid to the buccal base of the metaconid, and together with the protoconid, metalophid and metaconid isolate a small, narrow and crescent-shaped anterofossettid (Fig. 3A, E, M). The posterocingulid extends from the distobuccal aspect of the hypostylid, around the distobuccal corner of the tooth, and along the distal margin of the tooth for its entire width, ultimately connecting to the posterior arm of the entoconid. There is no trace of hypoconulid on the posterocingulid. The hypolophid and posterocingulid (clearly lower) delimit a narrow and buccolingually extended metafossettid, which appears as a distal gutter at the base of the gently sloping distal flanks of the hypoconid and hypolophid (Fig. 3A, E, G, J–K). Mesiolingually, the metaconid displays a long but low and arcuate posterior arm, which runs distally to reach a small and mesiodistally elongated mesostylid (Fig. 3E, G, K–L). The latter remains separated distally from the anterior arm of the entoconid, due to the presence of a wide V-shaped notch that punctuates the lingual crown margin, thereby maintaining the mesoflexid open lingually (Fig. 3K–L). There is no development of ectolophid in the centre of the talonid basin, so the mesoflexid is confluent with the hypoflexid, thus generating a central transverse furrow-like valley, open lingually and buccally (Fig. 3G–H, K–L). On either side of this transverse valley that bisects the tooth, irregularities on the enamel surface, appearing as roughnesses (or wrinkles), are present along the internal sloping flanks of the cuspids and cristids/lophids.
LACM 162449 m1 or m2. This lower molar is markedly larger than dp4 and subrectangular in occlusal outline (with rounded corners; Fig. 4A, E–F), being slightly wider than it is long (mesiodistal length: 1.98 mm, buccolingual width: 2.07 mm). Its occlusal surface is flat (Fig. 4K–P) due to an advanced stage of wear, which somewhat obscures the delineation of the main structures. The tooth exhibits a bilobed pattern with a narrow, deep and transverse central valley, and it is bounded mesially and distally by an anterocingulid and a posterocingulid, respectively (Fig. 4E–F, G, J). The two main transverse lophids are very wide, each displaying an enamel layer remarkably thick (either on their trailing edge or on their leading edge) as well as a narrow central dentine layer (Fig. 4A, D, E–F). The anterocingulid and posterocingulid are slightly affected by attritional wear (both cingulids being at the same occlusal level as that of the two main lophids, which are strongly abraded). Thus, considering the marked thickness of the transverse dentine layer that appears upon each lophid due to wear, it may be expected that the tooth, when pristine, was much more high-crowned, perhaps twice the height of the currently worn crown (Fig. 4G–H, K–L, N, P). The cusps/stylids are not clearly defined, as they are subsumed within the lophids, a condition which was likely achieved at an early stage of wear due to the strong thickness of the enamel layer. Despite the wear, the overall dental pattern of this molar (i.e. two internal lophids, peripheral cingulids, and a deep and transverse central valley bisecting the crown), is however basically similar to that of the LACM 162478 deciduous premolar (Fig. 3A–N). The bilophate pattern of LACM 162449 likely derives from the buccolingual coalescences of the bulbous and mesiodistally compressed lingual cusps with the similarly bulbous and mesiodistally compressed buccal cusps and stylids. Based on the pattern of the LACM 162478 dp4, on the LACM 162449 m1 or m2, the distal main lophid is recognized here as a hypolophid, which consists of a merged entoconid and hypoconid, plus a hypostylid that likely occupied a slightly more mesial position on the buccal cingulid. Likewise, the mesial main lophid is recognized as a metalophid, resulting from the fusion of the metaconid with the protoconid and a protostylid situated on the mesial part of the buccal cingulid. These two lophids show in their middle part some kind of mesiodistal pinching points, which could testify, to some extent, to the delineation of the regions of these main bulbous cusps (Fig. 4A, E–F, J). The anterocingulid stretches from the mesiolingual aspect of the metaconid to the anterobuccal corner of the tooth, the latter being assimilated to the mesial part of the buccal cingulid on which the protostylid of the metalophid would be included. Similarly, the posterocingulid
extends from the posterior arm of the entoconid, along the posterior margin of the tooth for its entire width, and around the distobuccal corner of the tooth, the latter being assimilated to the distal part of the buccal cingulid on which the hypostylid of the hypolophid would be included. The anterocingulid and posterocingulid are very closely positioned to the metalophid and hypolophid, respectively, but remain distinct from these two main lophids by the presence of very narrow and shallow transverse grooves (faintly visible), corresponding to the anterofossettid and metafossettid, respectively (Fig. 4A, E–F, G, J, M). The buccal and lingual margins of the crown are notched in their middle. These openings correspond to the extension of the central transverse and narrow valley that separates the two main lophids. This central valley is very deep with almost vertical sides (Fig. 4A, E–F, J), and corresponds to the confluence of a narrow hypoflexid with a similarly narrow mesoflexid.

Comparisons and discussion

The two specimens discovered represent two lower dental loci, documenting distinct ontogenetic stages, and testifying to the co-occurrence of both a new-born and an adult individual at LACM Loc. 8060. Despite such disparate morphological information, these two cheek teeth have a comparable overall Bauplan. The two teeth have a similar thickness of the crown enamel layer (i.e. very thick; Figs 3C–D, 4C–D), and their occlusal pattern is basically bilophodont, with a deep and transverse central valley bisecting the occlusal crown surface into two distinct main lophids, associated with peripheral cingulids (Figs 3, 4). These two dental specimens are further compatible in size and morphology, and are thus considered here as documenting a single rodent species. Understanding the occlusal morphology of the pristine deciduous premolar (dp4) was essential to decipher the structural pattern of the worn molar, especially to recognize plausibly homologous structures. Although autapomorphic in many features (see below), the occlusal pattern of these two lower teeth is distinguished primarily by the lingual and buccal cuspids plus unusual buccal stylar cuspids that are arranged in two well-separated transverse rows. Such a pattern of lower teeth characterized by the addition of stylids to lophids is autapomorphic for Geomyina among geomorph rodents (e.g. Wahlert 1983; Korth et al. 1991; Korth 1994; 2007; Flynn et al. 2008). This is particularly shown in lower teeth of North American Palaeogene geomyins (Fig. 5), such as heliscomyids (e.g. Heliscomys Cope, 1873), florentiamyids (e.g. Florentiamys Wood, 1936,
Kirkomys Wahlert, 1984), stem geomyoids (i.e. Geomyoidea incertae sedis sensu Flynn et al. 2008; e.g. Proharrymys Korth & Branciforte, 2007, Mojavamys Lindsay, 1972), heteromyids (e.g. Proheteromys Wood, 1932, Eochaetodipus Korth, 2008), and geomyids (e.g. Pleurolicus Cope, 1878, Gregorymys Wood, 1936), in which teeth are widened buccally by the addition of well-defined protostylid and hypostylid, positioned directly buccal to the protoconid and hypoconid, respectively. Analogous widenings also occur lingually on their upper teeth, by the addition of a prominent lingual cingulum bearing moderately to strongly developed styles. With abrasion and attrition wear, upper and lower teeth of Geomyina are well recognizable by their typical bilophate/bilobed pattern. The new West Indian taxon, Caribeomys, is here considered as being related to a Geomyina among Geomorpha due to its bilophodonty including buccal stylids in lophids. This condition does not match that of stem Geomorpha (Geomorpha incertae sedis sensu Flynn et al. 2008), such as representatives of the Oligocene-Miocene Jimomyidae (Jimomys Wahlert, 1976, Texomys Slaughter, 1981, and Zetamys Martin, 1974), or the Eocene Griphomys Wilson, 1940, Meliakrouniomys Harris & Wood, 1969, and Floresomys Fries et al., 1955, the geomorph status of which is questioned (see Flynn et al. 2008). Due to wear, the occlusal pattern of the lower molar of Caribeomys might, at first glance, be very close to that of lower molars of Meliakrouniomys (deep, narrow and buccolingually open transverse valley bisecting the tooth; short and very low anterocingulid and posterocingulid, closely spaced from, respectively, the thick transverse metalophid and hypolophid). However, the LACM 162449 molar displays a better defined posterocingulid, without hypoconulid, and a narrower transverse valley. The LACM 162478 premolar also displays a long and clearly better-defined posterocingulid, and above all it develops buccal stylids, which are absent in Meliakrouniomys. Lower teeth of Caribeomys differ from those of Griphomys in being slightly higher-crowned, in lacking the mesoconid, and in bearing a strong and complete hypolophid (absent in Griphomys). The dental pattern of Jimomys is only superficially reminiscent to that of Caribeomys. Indeed, at first sight, lower molars of Jimomys exhibit a bilobed pattern (with wear), but this is a “false” bilophodonty. The anterocingulid and posterocingulid of lower molars of Jimomys are not low, but particularly well developed (i.e. strong and elevated), and as such described as anterolophid and posterolophid, respectively. Despite the presence of a deep central transverse valley bisecting the crown mesiodistally in two separated lobes, as observed in Caribeomys, the bilobed dental pattern characterizing lower molars of Jimomys, derives from a tetralophodont pattern. Compared with Caribeomys, lower molars of Jimomys are higher-crowned, they do not show any development of buccal stylids, and the deciduous and permanent premolars are
pentalophodont, characterized by the addition of a strong central lophid (i.e. a mesolophid), in contrast to the condition seen in the LACM 162478 dp4 attributed to *Caribeomys*.

The diagnostic characters of the different families of Geomyina so far recognized are primarily cranial features (notably for differentiating early representatives among extinct heliscomyids and florentiamyids, and extant heteromyids and geomyids), although dental characteristics are also mentioned (e.g. Wahlert 1983, 1985, 1991; Korth *et al.* 1991; Korth 1994, 2008; Flynn *et al.* 2008; Jiménez-Hidalgo *et al.* 2018; Calede & Rasmussen 2020). With only two isolated lower teeth documenting this Oligocene rodent from Puerto Rico, despite their apparent geomyin affinities, the question remains as to whether this new taxon, *Caribeomys*, could belong to one of the defined families, or constitute a distinct basal branch within the parvorder. In all Palaeogene taxa of Geomyina (Fig. 5), the buccal stylar cuspids (protostylid and hypostylid) are clearly much smaller than the four main bulbous cuspids, and remain separated from the protoconid or metalophid (i.e. when protoconid and metaconid are merged) and the hypoconid or hypolophid (i.e. when hypoconid and entoconid are merged). It is only at some advanced stage of wear that these two stylids incorporate the metalophid and the hypolophid. On the pristine LACM 162478 dp4 attributed to *Caribeomys*, the main cuspids are moderately bulbous, even appearing as discreet tubercles within the two main lophids. Because of this quite weak development of the main cuspids, the buccal stylids may in contrast appear well marked and salient, but they still remain lower than the buccal main cusps. In *Caribeomys*, one of the differences from other geomyin taxa is the connection of the protostylid and hypostylid with the protoconid and hypoconid, respectively (Figs 3, 5B). In this taxon, the buccal stylids are therefore incorporated in the lophids, not due to an advanced wear, but due to the presence of thin and very short cristulids. The LACM 162449 m1 or m2 attributed to *Caribeomys* is particularly worn, and the protostylid and hypostylid are undifferentiated in the buccal region of the metalophid and hypolophid, respectively (Figs 4, 5A). The same is true for the four main cuspids, which are merged together without clear remnant of delineations, and entirely subsumed within these two main and massive lophids. LACM 162449 is noticeably worn, but clearly not at a very advanced stage of wear given the still high elevation of the mesial and distal walls of the deep central and transverse valley separating the two main lophids and bisecting the tooth (Fig. 4). On this molar, the lack of delineation of the main linked structures forming the two massive lophids is a consequence of the great thickness of the enamel layer, a condition that is somewhat unusual, if not unique among the earliest geomyin rodents (a thick enamel layer is observed in some more recent
geomyoid taxa, such as for instance *Mojavemys* Lindsay, 1972, *Phelosaccomys* Korth & Reynolds, 1994, or *Schizodontomys* Rensberger, 1973). In early Geomyina, the absence of delineation of the main bulbous cusp(-id)s and absorption of stylar cusp(-id)s in the loph(-id)s are observed only in old individuals with extreme abrasion and attrition wear of low-crowned cheek teeth (in heliscomyids, florentiamyids, and in some heteromyids), or in young individuals of more advanced and younger taxa that have moderately to very high-crowned teeth (other heteromyids and geomyids).

The two teeth documenting *Caribeomys* are also characterized by the presence of well-defined and buccolingually extended anterocingulid and posterocingulid, as mesiolingual and distolingual extensions of the buccal cingulid bearing the stylids. The presence of still well-developed antero- and postero- cingulids is primarily observed in heliscomyids (most species of *Heliscomys*; Fig. 5C–E; e.g. Korth et al. 1991; Korth 1994, 2008a; Korth & Branciforte 2007; Kelly 2009) and, to some extent, in some low-crowned species of stem geomyoids (e.g. *Proharrymys*; Korth & Branciforte 2007; Fig. 5L) and heteromyids (such as *Proheteromys*; e.g. Korth 2014; Korth & Samuels 2015). However, *Caribeomys* differs from heliscomyids in having teeth much less brachydont (i.e. mesodont) but also much less bunodont, without well-differentiated cuspids and stylids. It also differs from the aforementioned stem geomyoids and heteromyids in having lingual and buccal cuspid which, when merged (with wear), form transverse metalophid and hypolophid, not the V-shaped lophids (two chevrons) observed in *Proheteromys* and *Proharrymys* for instance (Fig. 5L). The presence of antero- and posterocingulids in lower teeth of *Caribeomys* also contrasts with lower teeth of florentiamyids (e.g. *Kirkomys*, *Florentiamys*; Fig. 5G–H, J) and some other species of heteromyids (e.g. *Eochaeotodipus*, *Perognathus*; Fig. 5K), in which these cingulids are faintly marked, buccolingually reduced to absent (e.g. Wahlert 1983; Korth 1994, 2008b; Korth & Branciforte 2007). In the two latter families, when present, the anterocingulid is more frequently observed, but developed to varying degrees. Geomyids in general (geomyines and entoptychines; Fig. 5M–N) do not show any development of these two cingulids (in taxa with either low- or high-crowned teeth). With the retention of such cingulids, *Caribeomys* as well as heliscomyids and few heteromyids, exhibit therefore a primitive geomyin condition, which is found in stem Geomorpha (*Geomorpha incertae sedis sensu* Flynn et al. 2008) and characteristic of the Eomyidae, the extinct sister group of Geomorpha (e.g. Flynn 2008). However, in some Geomorpha *incertae sedis*, notably the Jimomyidae (see above; e.g. *Jimomys*, *Texomys*) and in Eomyidae (Eomyinae, Apeomyinae, and Yoderimyinae: e.g.}
*Adjidaumo* Hay 1899, *Paradjidaumo* Burke 1934, *Apeomys* Fahlbusch 1968, *Apeomyoides* Smith et al., 2006, *Eomys* Schlosser 1884, *Yoderimys* Wood 1955), the upper and lower tooth patterns were not as simplified as in Geomyina (i.e., bilophodonty), as their antero- and postero-cingulids turned into anterolophid and posterolophid, respectively, thereby generating “multilophodont” patterns (tetralophodonty or even pentalophodonty).

**DISCUSSION**

Despite a limited fossil material, this small early Oligocene Puerto Rican rodent exhibits dental attributes indicating undisputed basal geomyin geomorph affinities. However, it also displays a suite of startling specializations (e.g. very thick enamel layer, cuspsids and stylids not well differentiated and interconnected) that makes this insular taxon new and distinct from known Palaeogene representatives of all recognized groups of Geomyina (i.e. Heliscomyidae, Florentiamyidae, Heteromyidae, Geomyidae, and Geomyoidea incertae sedis; see Flynn et al. 2008). The familial assignment of *Caribeomys* thus remains indeterminate to date as it requires further morphological support than current data allow. But considering its diagnostic autapomorphous dental traits compared with coeval stem geomyins, the possibility also exists that *Caribeomys* is the representative of a separate geomyin lineage (i.e. a distinct West Indian branch) having evolved in insular context. Regardless of the systematics issues, the presence of a rodent with North American affinities in lower Oligocene deposits of Puerto Rico also raises puzzling palaeobiogeographical questions. This discovery challenges the traditional view that all extinct and extant rodents endemic to the Greater and Lesser Antilles are proximately derived from South American sources (e.g. MacPhee & Iturralde-Vinent 1995, 2005; Woods et al. 2001; MacPhee 2005; Fabre et al. 2014; Brace et al. 2015; Courcelle et al. 2019; Marivaux et al. 2020). The early Oligocene site at Río Guatemala in San Sebastián, northwestern Puerto Rico (LACM Loc. 8060), records two distinct rodent groups of radically different geographical origins: two chinchilloid caviomorphs of South American source (Marivaux et al. 2020), and a geomyin geomorph of North American source (Fig. 6). When and how did these rodents reach this Caribbean island, or more generally the Greater Antilles?

The Greater Antilles islands, including Cuba, Hispaniola and Puerto Rico (but not Jamaica), share geological histories dating back to the Cretaceous–Palaeogene Great Antillean
arc system (Mann 2007; Pindell & Kennan 2009; Stanek et al. 2009; Boschman et al. 2014). However, the island complex seemingly did not become permanently subaerial until the whole belt was tectonically deformed, i.e. not earlier than the late middle to late Eocene (e.g. Khudoley & Meyerhoff 1971; Lewis et al. 1990; MacPhee & Iturralde-Vinent 1995; Iturralde-Vinent & MacPhee 1999; Graham 2003; Iturralde-Vinent 2006). The permanent land emersion certainly promoted survival of terrestrial animals that invaded the Greater Antilles from that time. It can be hypothesized that either land connections of South, Central (Chortis block) and North America with the Greater Antilles and/or overwater transports between these lands mediated these early rodent dispersals. Although an overwater dispersal from South America to the Greater Antilles cannot be ruled out, the GAARlandia hypothesis of a subaerial connection between both landmasses (via the emerged Aves Ridge; Fig. 6) during a sea level drop recorded around the Eocene–Oligocene transition (e.g. Miller et al. 2020), is often advocated to explain the origin of several terrestrial organisms (including the chinchilloid rodents) of South American sources on the Caribbean islands (e.g. MacPhee & Iturralde-Vinent 1995, 2005; Iturralde-Vinent & MacPhee 1999; MacPhee 2005; Iturralde-Vinent 2006; Alonso et al. 2012; Matos-Maraví et al. 2014; Dziki et al. 2015; Chamberland et al. 2018; Delsuc et al. 2019; Presslee et al. 2019; Tong et al. 2019; Blackburn et al. 2020; Crews & Esposito 2020; Marivaux et al. 2020; Philippon et al. 2020). What about the Palaeogene dispersal of geomorph rodents from North America to the Greater Antilles? This relative sea level drop (e.g. Miller et al. 2020) might also have contributed to that dispersal in having generated a potential subaerial land bridge between southern North America (and/or Central America) and the Greater Antilles. There is so far no geological support for such a land connection at that time, but such a sea level drop might have contributed to shorten inter-island distances, thus facilitating successful overwater dispersals for terrestrial organisms.

Jamaica, originally part of the Central American tectonic plate, has yielded a fossil rhinocerotoid of North American origin (Fig. 6) and the semi-terrestrial stem sirenian Pezosiren portelli in middle Eocene deposits (Domning et al. 1997; Domning 2001a), as well as iguanid and crocodyliform remains (Berg 1969; Domning & Clark 1993; Pregill 1999; Vélez-Juarbe & Brochu in press), thereby demonstrating the occupation of this Caribbean island by terrestrial taxa, during its partially subaerial and perhaps continuous connection with the mainland at that time (MacPhee et al. 1999). However, Jamaica subsequently remained partially or fully submerged until its re-emergence in the mid-Miocene, and may never have been directly connected to the Cainozoic Antillean arc (e.g. Buskirk 1985; Morgan 1993;
Therefore, the overwater dispersal hypothesis (from mainland southern North America or from Jamaica) would better explain the pattern of colonization of the Greater Antilles (at least Puerto Rico) by North American stem geomorph rodents. Bracketing the age of that colonization remains difficult given the limited palaeontological data so far assembled, which does not allow establishing a time-calibrated phylogenetic context. Modern geomorph families (Geomyoidea: Geomyidae and Heteromyidae) diverged and adaptively radiated during the Oligocene and the early Miocene (e.g. Hafner 1993; Hafner et al. 2007; Montgelard et al. 2008; Meredith et al. 2011; Tapaltsyan et al. 2015; Jiménez-Hidalgo et al. 2018), while their extinct close relatives (Heliscomyidae, Florentiamyidae, and both stem Geomyoidea and few Geomyidae) are well represented as early as the late Eocene and early Oligocene in the fossil record of Central and North America (e.g. Korth 1994; Korth & Eaton 2004; Flynn et al. 2008; Kelly 2009; Jiménez-Hidalgo et al. 2015, 2018; Ortiz-Caballero et al. 2020). The most likely ancestry of Geomorpha (as well as their sister groups Eomyidae and Castoroidea; e.g. Flynn 2008; Flynn and Jacob 2008) is expected from a middle Eocene North American sciuravid stock (e.g. Wood 1935; Wilson 1949; Wahler 1985, 1991; Korth 1994; Marivaux et al. 2004; Flynn et al. 2008; Walton & Porter 2008). The peculiar dental specialization characterizing Caribeomys with respect to its coeval North American stem geomyin counterparts may indicate that its lineage had diverged for some time. This would thus suggest that the dispersal of geomorphs towards the West Indies likely occurred well before rather than shortly before its appearance in the Puerto Rican fossil record, i.e. sometime during the late Eocene–earliest Oligocene interval.

Interestingly, this Puerto Rican geomorph rodent adds to the very few cases of Antillean terrestrial vertebrates of North American origins. The recent discovery of the plethodontid salamander *Palaeopleodon hispaniolicae* in ancient amber-bearing deposits (early-middle Miocene) of the Dominican Republic, represents a taxon of North and Central American affinities with no extant relatives in the region (Iturralde-Vinent & MacPhee 1996; Poinar & Wake 2015). The most famous example are the solenodonotans (Eulipotyphla), which are widely believed to have arrived in the Greater Antilles during the late Cretaceous, with a Palaeogene divergence of the two known West Indian families, the recently-extinct Nesophontidae and the extant Solenodontidae (Roca et al. 2004; Brace et al. 2016; Springer et al. 2018). The timing of solenodonotan colonization would be consistent with
palaeogeographical reconstructions that place the proto-Greater Antilles close to North America (Iturralde-Vinent & MacPhee 1999; Roca et al. 2004; Iturralde-Vinent 2006), in a period otherwise too early to explain the origin of the Antillean geomorph rodent. However, as the island complex of the Antillean arc system was not permanently subaerial before the late middle/late Eocene (e.g. Khudoley & Meyerhoff 1971; Lewis et al. 1990; MacPhee & Iturralde-Vinent 1995; Iturralde-Vinent & MacPhee 1999; Graham 2003; Iturralde-Vinent 2006), then the possibility also exists that solenodonotans derive from a dispersal during the mid-Palaeogene of a now-extinct Palaeogene North American soricomorph relative (Asher et al. 2002; Roca et al. 2004; Brace et al. 2016; Sato et al. 2016). In that context, the solenodonotan ancestor may have had a biogeographical history similar to that of the Puerto Rican geomorph rodent. In the absence of clear subaerial land connections between southern North America and/or Central America with the Greater Antillean arc after the Palaeocene, overwater dispersals shall be advocated as a hypothesis for explaining such a colonization. By mid-Palaeogene times, the ancestors of the geomorph rodent and the plethodontid salamander – and possibly also the ancestor of solenodonotan eulipotyphlans – may have colonized the Greater Antilles, either through a single episode or a series of rafting episodes.

CONCLUSIONS

Without these two tiny teeth recovered from a single and limited spot in the Río Guatemala of Puerto Rico, it could not have been predicted that a stem geomorph of North American affinities was living in the Greater Antilles during the early Oligocene, alongside early chinchilloids of South American origins. Caribeomys is an ancient relative of modern North and Central American heteromyids (kangaroo rats and pocket mice) and geomyids (pocket gophers) and so far, it is the only evidence of the past presence of geomorphs in the West Indies. This rodent group is indeed absent from modern and Quaternary West Indian ecosystems, and it has never been reported in the rare Miocene records of the Greater Antilles or in the Eocene deposits of Jamaica (e.g. MacPhee et al. 2003; Blackburn et al. 2020). Morphological comparisons of Caribeomys with its Palaeogene North American counterparts indicate that this island lineage had diverged for some time, and likely dispersed toward the West Indies long before its appearance in the Puerto Rican fossil record (i.e. during the late Eocene–earliest Oligocene interval). What specifically were the modalities (overland versus...
overwater dispersal) and pathways for this rodent to reach Puerto Rico, or more generally the Greater Antilles? How long did this group survive in the Greater Antilles, at least in Puerto Rico? When did it become extinct? What caused its extinction? Geomorphs have flourished from the Oligocene to the present-day in North and Central America, but may only have been transitory in some Caribbean islands. Only fossil data will shed light on these issues. Early chinchilloids also documented in the Río Guatemala are possibly related to the Pleistocene–Holocene West Indian “giant chinchilloid hutias” (see Marivaux et al. 2020). The latter were only recently extirpated from the West Indies (MacPhee 2009), where they possibly evolved in isolation for over 30 million years. Among solenodonotans, only Solenodontidae still live in the West Indies today, but they are endangered, occupying a relictual distribution in Cuba and Hispaniola (e.g. MacPhee 2009; Turvey & Fritz 2011; Brace et al. 2016; Sato et al. 2016; Brandt et al. 2017). Its sister group, the Nesophontidae, was also extirpated recently (MacPhee et al. 1999). A similar pattern is observed across other terrestrial vertebrates, some with a long evolutionary history in the region (e.g. sloths, primates), which then became extinct over the last several thousand to hundreds of years due to climatic and/or anthropogenic causes (e.g. MacPhee 2009; Kemp & Hadly 2015; Cooke et al. 2017).

The LACM Loc. 8060 fossil-bearing locality provides only a glimpse into the palaeodiversity of rodents in the Greater Antilles, at least in Puerto Rico, some 29 million years ago. Such a heteroclite rodent palaeodiversity was somewhat unexpected, and it is probably far to be exhaustively documented. However, it offers new and invaluable insights into rodent palaeodiversity, and reveals how severely under-sampled palaeontologically the West Indies are, notably for deep times. Improving our knowledge of the past diversity of West Indian vertebrates is urgently needed to further our understanding of their evolutionary history, and to dedicate efforts towards their conservation.

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DATA ARCHIVING STATEMENT

This published work and the nomenclatural acts it contains, have been registered in ZooBank: http://zoobank.org/References/8C2E090D-E1EC-4269-9497-E2888B26F279. The raw files for the μCT-scan data and renderings of the 3D digital models of the fossil dental specimens described in this article are available on the online open-access platform MorphoMuseuM: https://doi.org/10.18563/journal.m3.128.

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**FIG 1.** Location and geological maps of the LACM Loc. 8060 fossil-bearing locality. A, the locality (yellow star) is located in the Río Guatemala, which flows through the township of San Sebastián, Puerto Rico (West Indies). LACM Loc. 8060 occurs in shallow marine deposits corresponding to the lower part of the San Sebastian Formation, which outcrops along the banks of the Río Guatemala (Vélez-Juarbe et al. 2007, 2014; Marivaux et al. 2020). $^{87}$Sr/$^{86}$Sr mean ages-dates calculated on low-Mg calcite shells of *Kuphus incrassatus* bivalves allowed for bracketing the age of the LACM Loc. 8060 between 29.78 and 29.17 Ma (late Rupelian) (Ortega-Ariza et al. 2015). B, panoramic view of LACM Loc. 8060, showing our team excavating the 15–20 cm thick layer of grey silty claystone (white dotted line) that yielded the fossil remains of mammals (photograph by J. Vélez-Juarbe).
FIG 2. Dental terminology (updated and modified after Wahlert 1983; Korth 1994, 2007; Marivaux et al. 2004; Calede & Glusman 2017; Calede & Rasmussen 2020). The terminology is applied on the two fossil teeth from LACM Loc. 8060: A, LACM 162449, right lower molar (m1 or m2); B, LACM 162478, right fourth lower deciduous premolar (dp4). Abbreviations: Aaet, anterior arm of the entoconid; Acd, anterocingulid (or anterolophid); Afxd, anteroflexid or paraflexid (here anterofossettid); Bcd, buccal cingulid; Et, entoconid (or entoconid region); Hd, hypoconid (or hypoconid region on the molar); Hfxd, hypoflexid; Hld, hypolophid (= the main distal lophid resulting from the coalescence of the bulbous entoconid with the bulbous hypoconid + the hypostylid); Hyd, hypostylid; Md, metaconid (or metaconid region on the molar); Med, metalophid (= main mesial lophid resulting from the coalescence of the bulbous metaconid with the bulbous protoconid + the protostylid on the molar); Msfxd, mesoflexid; Mstd, mesostylid; Mtfxd, metaflexid (here metafossettid); Pamd, posterior arm of the metaconid; Prd, protoconid (or protoconid region on the molar); Pryd, protostylid; Psd, posterocingulid (or posterolophid). The white arrow indicates the orientation of the teeth on the jaw (mesiolingual). The images of the teeth are renderings of 3D digital models of the fossil specimen, obtained by X-ray µCT surface reconstruction. Scale bar represents 1 mm.
FIG 3. *Caribeomys merzeraudi* gen. et sp. nov. from the late early Oligocene of Puerto Rico (LACM Loc. 8060). A–N, LACM 162478, right fourth lower deciduous premolar (dp4) in occlusal (A and E-E’ stereo pair), ventral (F), lingual (B, G–H), distal (I–J), buccal (K–L) and mesial (M–N) views. Sagittal cross-sections (CS) were performed (B) for visualizing enamel thickness at different crown heights. CS1 was made near the cervix (C) and CS2 near the occlusal surface (D). The occlusal view in A (whitened using magnesium smoke) is the result of the fusion of multi-focus images. The images in B and E–N are renderings of 3D digital models of the fossil specimen, obtained by X-ray µCT surface reconstruction. The images in C and D are 2D orthoslices of the specimen. Scale bar represents 1 mm.
FIG 4. *Caribeomys merzeraudi* gen. et sp. nov. from the late early Oligocene of Puerto Rico (LACM Loc. 8060). A–P, LACM 162449, right lower molar (m1 or m2) in occlusal (A, E–F), lingual (B, G–H), ventral (I), buccal (J–L), mesial (M–N) and distal (O–P) views. Sagittal cross-sections (CS) were performed (B) for visualizing the enamel thickness at different crown heights. CS1 was made near the cervix (C) and CS2 near the occlusal surface (D). The occlusal view in A is the result of the fusion of multi-focus images. The images in B and E–P are renderings of 3D digital models of the fossil specimen, obtained by X-ray µCT surface reconstruction. The images in C and D are 2D orthoslices of the specimen. Scale bar represents 1 mm.
FIG 5. Lower dentition of some early geomyin geomorphs for comparisons with lower teeth of *Caribeomys merzeraudi* gen. et sp. nov. from the early Oligocene of Puerto Rico. A, LACM 162449, right m1 or m2 of *C. merzeraudi*; B, LACM 162478, right dp4 of *C. merzeraudi*; C, UNSM 425737, right dp4–m3 (here dp4–m2) of *Heliscomys ootranderi* from the early Oligocene of Wyoming, USA (Korth et al. 1991, fig. 4E); D, UNSM 130501, right p4–m2 of *Heliscomys* sp. from the late Oligocene of Nebraska, USA (Korth & Branciforte 2007, fig. 2D); E, UMPE 0116, left dp4–m1 (reversed) of a geomyid indet. from the late Eocene of Oaxaca, Mexico (Jiménez-Hidalgo et al. 2015, fig. 4B); F, UNSM 24099, 24102 and 24102, left dp4 (reversed) and right m1 or m2 of *Sanctimus* cf. *S. stuartae* from the early
Miocene of Nebraska, USA (Korth 1992, fig. 14I, K–L); G, MCZ 5051, left p4–m3 (reversed) of *Kirkomys nebraskensis* from the late Oligocene of Nebraska, USA (Korth & Branciforte 2007, fig. 4B); H, SDSM 57391, left dp4–m2 (reversed) of *Kirkomys martintau* from the middle Oligocene of South Dakota, USA (Korth, 2014, fig. 9D); I, UNSM 24103, left p4–m3 (reversed) of *Hitonkala macdonaldtau* from the early Miocene of Nebraska, USA (Korth 1992, fig. 13C); J, Amherst 27-126, right p4–m3 of *Florentiamys loomisi* from the early Miocene of Wyoming, USA (Wood 1936, fig. 1–5; Stehlin & Schaub 1951, fig. 526); K, UNSM 130000, right p4–m3 of *Eochaeotodipus asulcatus* from the early Miocene of Nebraska, USA (Korth 2008b, fig. 2B); L, UNSM 130217, left p4–m3 (reversed) of *Proharrymys wahlertii* from the late Oligocene of Nebraska, USA (Korth & Branciforte 2007, fig. 6C); M, F:AM 51309, left dp4–m2 (reversed) of *Pleurolicus cf. sulcifrons* from the latest Oligocene of New Mexico, USA (Gawne 1975, fig. 9A); N, UMPE 622, right p4–m3 of *Gregorymys veloxikua* from the late Eocene of Oaxaca, Mexico (Jiménez-Hidalgo et al. 2018, fig. 7d). Scale bars represent 1 mm. The scale bar at the top centre of the figure is for A–M. The scale bar at the bottom right of the figure is only for N.
FIG 6. Palaeogeographical map of the Caribbean region by the Eocene–Oligocene Transition (34 Ma) and mammalian assemblages of interest (Eocene–Oligocene). Light landmass colours stand for areas submerged under shallow water. C, Cuba; E, early; Eoc, Eocene; GrANoLA, Greater Antilles-Northern Lesser Antilles landmass; H, Hispaniola; J, Jamaica; L, late; Mid, middle; Oli, Oligocene; PR, Puerto Rico. This original figure is based on data from Patton (1969), Domning (2001a, b), Iturralde-Vinent (2006), Jiménez-Hidalgo et al. (2018), Young et al. (2019), Cornée et al. (2020, 2021), Marivaux et al. (2020), and Philippon et al. (2020), and deeptimemaps.com.