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1 Multispecimen domain-state corrected paleointensity determination of a
2 detailed single lava flow, Xitle volcano (Mexico)

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17 **Abstract**

18 Determining paleomagnetic field intensity (paleointensity: PI) for lavas with high reliability
19 and low measurement uncertainty is still difficult to achieve. In addition to the factors on
20 which the PI used methods depend, this could be attributed to some non-ideal physical and
21 magnetic characteristics of lava sample, including grain size, cooling rate effect, and thermal
22 stability. Xitle volcano (SW Mexico City) is a good example to illustrate and discuss this
23 problem because dozens of previous PI studies were carried out on its evolved flow units,
24 which have commonly resulted in different mean values with large dispersions. Indeed, 211
25 published PI data obtained by use of Thellier and microwave experiments gave a mean of
26 $64.1 \mu T$ with a standard deviation of $11.0 \mu T$. After a careful evaluation, we found that only
27 134 of these data can be considered reliable, as they meet a set of selection criteria designed

28 in this study. These evaluated data gave an average mean of $62.0 \pm 9.3 \mu T$. In order to
29 strengthen the PI estimates of Xitle, we conducted a multispecimen domain-state corrected
30 (MSP-DSC) method along one vertical ($\sim 4.5\text{m}$) and three horizontal ($\sim 1.25\text{m}$, each)
31 profiles. Top horizontal and vertical profiles have fulfilled a stringent criteria set while central
32 and bottom profiles exceeded the alteration check criteria limit and thus are considered
33 unreliable. Accordingly, Xitle PI mean derived from MSP-DSC experiment is calculated at
34 $60.5 \pm 4 \mu T$, thus in a good agreement with the mean value estimated from previous filtered
35 data. The result and success rate obtained may be ascribed to cooling rate variations
36 commonly found at the lava profile, and indicate that MSP-DSC outcome is governed by the
37 magnetic properties such as the domain-size behavior and the thermal stability of the
38 magnetic carriers present in the treated specimens, as in the conventional Thellier &
39 microwave-style experiments. From these two averages, a combined mean and standard
40 deviation of $61.9 \pm 9 \mu T$ is calculated, which technically is considered the most probable
41 intensity estimate at the Xitle eruption time, *ca.* 370 AD.

42 Keywords: paleointensity; lavas; rock magnetic properties; multispecimen method; Xitle;
43 Mexico

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47 **1. Introduction**

48 Over the long history of our planet, the magnetic field generated in the liquid outer core
49 changed on different time scales from years to billions of years. Understanding the spatio-
50 temporal evolution of this field requires careful determination of its strength. Because of its
51 large contributions in deciphering the geodynamo behavior (Biggin et al., 2012) and
52 improving the global geomagnetic field models (e.g., SHA.DIF.14k, Pavón-Carrasco et al.,
53 2014), several methods were proposed to obtain a reliable estimate of the PI: the classical

54 Thellier-Thellier experiment (Thellier & Thellier 1959) and other protocols (e.g., Coe et al.,
55 1967; Aitken et al., 1988; Tauxe and Staudigel, 2004); the Shaw method (Shaw, 1974) and
56 its variants (e.g. Tsuanakawa and Shaw, 1994); pseudo-Thellier (Tauxe et al., 1995); the
57 microwave technique (Hill & Shaw 1999); and the recent approach of multispecimen
58 (Biggins and Poidras, 2006; Dekkers and Böhnel, 2006; Fabian and Leonhardt, 2010).
59 Despite the improvements achieved in the laboratory protocols, these methods give reliable
60 intensity with a low success rate (generally less than 30%) from basalts, the material of
61 interest in this study. This was showed by comparing PI data retrieved from historically
62 erupted lava flows, such as those in Hawaii (e.g. Yamamoto et al., 2003; Böhnel et al., 2011;
63 Grappone et al., 2019) and Etna (e.g. Hill and Shaw, 1999; De Groot et al., 2013), with the
64 actual geomagnetic field intensity that is well known from geomagnetic observatories. Th
65 elow success rate in the PI methods to recover the expected field intensity with high accuracy
66 could be attributed to several reasons, including the presence of non-ideal physical and
67 magnetic properties, magneto-mineralogical alteration, the cooling rate difference and
68 presence of local magnetic field effects (Stacey & Banerjee 1974). From the PI point of view,
69 lava samples must contain ferromagnetic particles of single domain (SD; $< \sim 80$ nm) to
70 pseudo-SD (PSD; $< \sim 0.1$ μm) size . This condition however, is not easily reached because in
71 naturally cooled lavas there are always contributions from grains larger than ~ 1.0 μm , of
72 multidomain (MD) size. In this context, Cromwell et al. (2015) have showed the capability
73 of subaerial basaltic volcanic glass to give accurate field intensity as they have cooled quickly
74 and behaves as SD particles. Unfortunately, these glassy samples are not usually available,
75 but indeed the most commonly encountered material is a lava flow which depends on its

76 position and cooling rate can take from days to several months to cool, and thus a wide range
77 of domain size is expected.

78 All these factors may be responsible for over- or under-estimating the PI values. Xitle lava
79 flows (Fig. 1) allowed to illustrate and discuss this problem because dozens of previous PI
80 estimates conducted on them have commonly yielded different mean values with large
81 dispersions . These data were mainly obtained by means of Thellier method and some others
82 were provided by means of microwave and Shaw techniques. It should be mentioned here
83 that previous data are of uneven quality and thus in the next section we will discuss their
84 reliability based on today's set criteria parameters (Paterson et al., 2014; Thellier et al., 2014).
85 The present work was designed to reinforce Xitle PI estimates and to reduce its errors through
86 applying MSP-DSC method along one vertical (*ca.* 4.5m) and three horizontal profiles.
87 Providing a new reliable PI value for a dated flow unit is important to (among others) enhance
88 the global harmonic spherical models of the secular variation of the geomagnetic field over
89 the last millennia (e.g. Nilsson et al., 2014 and Pavón-Carrasco et al., 2014) and improve our
90 knowledge of the local field intensity behavior in central Mexico at *ca.* 370 AD.

91 **2. Geological setting and sampling**

92 The study profile (19.328° N; 99.189° W) is part of the flow VI of Xitle (Fig. 1) located
93 inside the campus of UNAM (Universidad Nacional Autónoma de Mexico). Xitle lies in the
94 central sector of the Trans-Mexican Volcanic Belt (TMVB, Fig. 1) which is an E–W trending
95 zone of *ca.*1000 km length extending from the Pacific Ocean to the Gulf of Mexico. The
96 TMVB is divided in western, central, and eastern sectors. The Sierra de Chichinautzin
97 Volcanic Field (SCVF) is located in the central sector (inset Fig. 1).

98 The SCVF contains high concentration of monogenetic volcanoes with about 220 Quaternary

99 volcanic products including cinder cones and lava flows (Siebe, 2000; Rodríguez-Trejo et
100 al., 2019) of wide compositional range. Xitle is considered as the youngest monogenetic
101 volcano of the SCVF with a radiocarbon age of 1530-1630 uncalibrated yr BP (cal. 370 ± 60
102 AD; Siebe et al., 2004; Arce et al., 2013). Xitle eruption is an example of the impact of
103 volcanic disaster on the human population, as supposedly it had damaged the pre-Hispanic
104 settlements around Cuicuilco pyramid (Fig. 1) which prompted them to emigrate (Siebe,
105 2000), as like in the historical eruptions of the Jorullo (1759–1774 AD; Guilbaud et al., 2011;
106 Rasoazanamparany et al., 2016; Alva-Valdivia et al., 2019) and Paricutin (1943–1952 AD;
107 Luhr and Simkin, 1993; Pioliet al., 2008) monogenetic volcanoes.

108 The sampled site was selected so that the bottom and top of the lava section (Fig. 3) are
109 visible. Sampling was done using a portable gasoline powered drill, and 72 core samples,
110 each with 5-10 cm long and 2.5 cm diameter, were collected. In this study, four profiles (Fig.
111 3) were taken and distributed as follow: one vertical profile (V) of *ca.* 4.5 m thickness and
112 composed of 43 cores; and three horizontal profiles (H) of *ca.* 1.25 m length for each: top
113 horizontal (HT) of 11 cores; middle horizontal (HM) of 10 cores; and the bottom horizontal
114 (HB) of 11 cores.

115 **3. Previous PI studies**

116 Nine PI studies were conducted on Xitle lava flows by means of the double heating Thellier
117 (Nagata et al., 1965; Urrutia-Fucugauchi, 1996; Gonzales et al., 1997; Alva Valdivia, 2005;
118 Böhnel et al., 1997; Morales et al., 2001, 2006; Mahgoub et al., 2019); microwave (Böhnel
119 et al., 2003); and Shaw (Urrutia-Fucugauchi, 1996; Gonzales et al., 1997) methods. Sampling
120 in five of these studies were collected randomly and its coordinates are of low precision
121 which makes us unable to define target flow unit in some of them. On the other hand, three

122 studies (Böhnel et al. , 1997; 2003; Alva-Valdivia et al., 2005) were designed so as to sample
123 vertical profiles over a specific cooling unit.

124 Two studies (Böhnel et al., 2003; Mahgoub et al., 2019) were carried out on pottery fragments
125 that were reheated by Xitle eruption and thus acquired their magnetization at the same time.
126 Böhnel et al. (2003) have performed microwave experiments on lavas and pottery fragments,
127 with the field applied perpendicular and parallel to the their NRMs. We note also that PI data
128 points presented in Böhnel et al. (1997) have been re-analyzed by Böhnel et al. (2003)
129 applying a stringent set of selection criteria. At this point, it must be stated that Böhnel et al.
130 (1997) carried out their study rather to find out how PI varies over the Xitle flow (see section
131 5.2) and if there is a relation between rock magnetic properties and success rate, therefore
132 they have not used a very strict selection criteria. Morales et al. (2006) tried to figure out the
133 cause of PI dispersion through conducting cooling rate correction. Based on their results, a
134 significant decrease in PI-dispersion was obtained (from 7.5 to 3.5 μT), thus they have
135 claimed that cooling-rate effect may have a prominent role in the observed dispersion. There
136 are two points to be mentioned in this context: the first is that Morales et al. (2006) have
137 obtained positive and negative corrections from nearby samples (of 10-20 cm distance) that
138 should have a very similar cooling rate. Secondly, if the change in the cooling rate is the
139 reason for the PI-dispersion, then, reasonably, Thellier results will give less scatter than
140 microwave approach, as the duration of each microwave step is *ca.* 10 seconds (Hill and
141 Shaw, 1999). Böhnel et al. (2003), though using microwave technique on a nearby profile,
142 provided PI mean results with similar dispersion as commonly obtained in Thellier. These
143 two points most likely rule out the effect of cooling rate as a major cause of intensity variation
144 acquired from Xitle volcano.

145 In order to demonstrate the PI mean and scatter in each study and to illustrate the consistency
146 between different studies, we plot the PI data at specimen level on Figure 2. They are 214
147 data points [211 derived from Thellier and microwave and 3 from Shaw method] considered
148 acceptable by the author(s) of each work. As Figure 2 illustrates, the data from each study is
149 highly scattered and the whole data are inconsistent as well. We note that the PI-dispersion
150 of these studies cannot be compared as the number of data is not equal: half of data were
151 obtained by Böhnelt et al. (1997) and Alva-Valdivia et al. (2005) while Urrutia-Fucugauchi
152 (1996) and Gonzalez et al. (1997) provided only 3 PI data points using Shaw technique. Using
153 all data, we calculate Xitle PI mean at 64.1 μT (Fig. 2) and the 95% standard deviation (σ) is
154 11 μT , *ca.* 18% of the mean. Since the number of data is generously available and meet the
155 suggestion of Biggin et al. (2003) that PI mean value of any lava flow should be based on as
156 many samples as possible (at least five). Apparently this dispersion could be related to,
157 among others, some problems in the experimental PI methods (especially those applied long
158 time ago). We evaluate previous 211 Thellier and microwave - derived PI data in terms of
159 recently proposed reliability standards (e.g. Tauxe and Staudigel 2004; Chauvin et al., 2005;
160 Paterson et al., 2014). There are some other problems that could be seriously responsible for
161 such dispersion, including alteration of the ferromagnetic particles, cooling rate effect, and
162 presence of local magnetic field effects. These effects will not be addressed here.

163 Besides the low number of provided data, the two studies that used Shaw experiment
164 (Urrutia-Fucugauchi, 1996; Gonzalez et al., 1997) cannot be considered reliable as they did
165 not perform alteration tests (e.g. Tsunakawa and Shaw, 1994). In Thellier experiments,
166 samples are heated up gradually from low (e.g. 100 °C) to high temperature (commonly
167 below curie point) and the natural remanent magnetization (NRM) is consecutively replaced

168 by laboratory induced thermal remanent magnetization (TRM), in a known laboratory field.
169 This must be done with some alteration checks (Coe et al., 1978) in order to ensure that no
170 alteration occurred during repeated heating. Laboratory procedure of microwave method is
171 the same as Thellier but instead of heating in a conventional oven, samples are demagnetized
172 by exposure to high-frequency microwave (Walton et al., 1993). It has been proved by Hill
173 et al. (2002) that microwave demagnetization is equivalent to thermal counterparts implying
174 that sample's NRM is replaced by a laboratory induced TRM, for further details we refer to
175 the work of Hill and Shaw (2000) and Böhnel et al. (2003). Due to similarity in experimental
176 approaches, we have set for PI data derived from Thellier and microwave the same criteria
177 set, which are:

- 178 1) Treated sample must have been checked for thermal alteration during heating by means of
179 the pTRM check criterions (δCK and/or δpal);
- 180 2) The stability of the sample's NRM directions during the experiments must have been
181 evaluated by one or all next parameters: MADanc, α , or DANG;
- 182 3) Following Biggin et al. (2003) suggestion, at least 5 specimens must have been used to
183 compute lava flow mean intensity with $\sigma \leq 10 \mu T$ or $\leq 20\%$ of the mean.

184 Applying these criteria set, we found that data presented in the work of Nagata et al. (1965),
185 Urrutia-Fucugauchi (1996), Gonzalez et al. (1997), Böhnel et al. (1997), Morales et al. (2001)
186 do not satisfy the mentioned criteria (Fig. 2). On the other hand, studies of Böhnel et al.
187 (2003), Alva Valdivia (2005), Morales et al. (2006), and Mahgoub et al. (2019) are reliable.
188 We calculate the Xitle mean PI from the 134 data considered as reliable data at $62.0 \mu T$ with
189 σ of $9.3 \mu T$ (see Fig. 2). This mean value is slightly below the mean calculated from all data

190 and the error is also reduced, however statistically they are indistinguishable at the 95%
191 confidence limit.

192 **4. Methods**

193 Rock magnetic experiments represented by the susceptibility versus temperature (k -T)
194 analyses and hysteresis measurements were done on 2-3 samples from each sampled profile
195 in order to check the magnetic variability in both horizontal and vertical directions.
196 Alternating field (AFD) and thermal demagnetization (THD) were measured in all the
197 samples of each profile. k -T curves were carried out up to $\sim 700^\circ\text{C}$ with a Bartington-MS2
198 susceptibility-meter coupled with the furnace XXXX (????), and the mean Curie temperature
199 (T_c) were defined as the inflection point after peaks in k . We evaluate the thermal alteration
200 that could occur during the laboratory heating by calculating a reversibility
201 parameter: $RP\% = \frac{k_h - k_c}{k_h} * 100$; where k_h and k_c represent values of k at heating and cooling
202 curves at 100°C , respectively (reference ???). Zero RP% indicates that the heated specimen
203 does not experience alteration. Hysteresis analyses were executed with the AGFM
204 (Alternating Gradient Force Magnetometer) for which samples weighting 5 to 40 mg were
205 used. The hysteresis parameters (saturation magnetization (M_s); saturation remanent
206 magnetization (M_{rs}); coercive force (H_c); and coercivity of remanence (H_{cr})) lead to have
207 an idea of the magnetic domain state of the magnetic carriers, Day diagram (Day et al., 1977).
208 AFD measurements were progressively applied from 5 to 100 mT with an AGICO LDA-3
209 equipment. Also, THD was carried out with ASC TD-48 thermal demagnetizer model in
210 every 50°C from 100 to 500°C and then from 530 to 600°C in 30°C step. From the
211 demagnetization measurements, we have calculated the median destructive field (MDF) and
212 median destructive temperature (MDT), which defined how the alternating field and

213 temperature values of the NRM loses half of its value.

214 PI experiments were estimated with the multispecimen method (Dekkers and Böhnel, 2006)

215 through which specimens are heated only once in different DC fields directed, independently,

216 parallel to the their NRM. The original protocol includes two steps: m0 and m1. Thereafter,

217 two additional steps (m2 and m3) were proposed by Fabian and Leonhardt (2010) in order to

218 correct for the domain state effect, and one more step (m4), where we repeat m1, is proposed

219 to check for any mineralogical alteration occurred during the experiment. In this study, 37

220 specimens were taken from all profiles to conduct the original (MSP-DB; referred to Dekkers

221 and Böhnel) and corrected (MSP-DSC; referred to domain state correction) protocols.

222 Heatings were carried out by use of a new infra-red-heating ultra-fast furnace developed and

223 available in the Geosciences Montpellier laboratory, called 'FURemAG'. The heating-

224 cooling time in the FURemAG furnace lasts 45 minutes. Based on the *k*-T curves and THD

225 results, the set-temperature will be selected so as to ensure unblocking sufficient portion of

226 NRM. Thus, we can get steeped linear fit with small confidence limit (Monster et al., 2015a),

227 and also magneto-mineralogical alteration can be avoided. To eliminate unwanted viscous

228 component in the NRM, the specimens were heated to 100°C and cooled to room temperature

229 in zero field, before the NRM measurement (m0 step). To be consistent with this pre-

230 treatment, the low temperature pTRM[100°C, Troom] was removed in the same way after

231 each pTRM acquisition (m1, m2, m3, and m4 steps) involved in the MSP-DSC protocol. The

232 magnetic remanence was measured with a cryogenic magnetometer (2G).

233 **5. Results**

234 ***5.1 Magnetic properties***

235 The *k*-T curves (Figs.4a) indicate the presence of several magnetic minerals in distinct

236 proportion. The Tc range from 540 to > 600 °C and RP% from 4 to 80%, suggesting the

237 presence of magnetite (Mag), Ti-poor titanomagnetite (Ti-poor TMag), and hematite with
238 varied reversibility degrees. In all studied samples, k value decrease after heating, indicating
239 that enclosed magnetic minerals have been oxidized. We found that both T_c and RP% do not
240 have any systematic behavior vertically or horizontally, but we can mention that HT and HM
241 profiles have moderate to good reversibility, respectively. Specimens of HB gave dissimilar
242 results and thus no clear conclusion can be outlined.

243 Hysteresis analyses show that all investigated samples are located in the range of PSD field
244 (Fig. 4b), which may suggest presence of a mixture of SD and MD particles in different
245 percentages. HT samples are located close together in the Day plot while those of HM and
246 HB did not show similar consistency. We deduced from these observations that HT profile
247 has a small homogenous PSD particles while, on the other hand, the middle and bottom
248 profiles have somewhat larger ferromagnetic particles of widespread type and/or size.
249 Apparently, few samples are provided this explanation and thus more samples would lead to
250 track better the domain state along vertical and horizontal profiles. However, the recent
251 findings of Roberts et al. (2018) should be mentioned where they find out that domain state
252 of a sample cannot be grasped simply from the Day plot as the hysteresis parameters are
253 based on several variables (Roberts et al., 2018).

254 Intensity-decay curves along the three horizontal profiles results from AFD measurements
255 (Fig. 4c) do not show any systematic behavior (as in previous experiments), and the MDF
256 ranges from 5 to 55 mT thus indicating the presence of varying magnetic grain composition
257 along the three sampled profiles. The THD data (Fig. 4d) showed the appearance of common
258 T_c point of magnetite (*ca.* 560°C) with small contribution from hematite, thus in agreement
259 with k -T curves. The MDTs are from 300 to 500 °C with a tendency of HB's samples (58,

260 59 and 64) to have lower MDT in comparison to HT and HM. This tendency could be
261 attributed to the presence of different Ti contents (as they have low unblocking temperature
262 spectra) or to large magnetic minerals size in the HB samples.

263 To sum up, magnetic experiments showed that rock from the sampled profiles, although of
264 limited vertical and horizontal spread, have wide range of type composition and size of the
265 enclosed magnetic minerals.

266 *5.2 MSP-DSC results*

267 The multispecimen data corrected for domain state (MSP-DSC) were analyzed with MSP-
268 Tool (Monster et al., 2015) software. The set-temperature throughout the experiments is
269 400°C and applied DC fields range from 10 to 80 μT . The domain state proxy (α -parameter)
270 was set to 0.5, as proposed by Fabian and Leonhardt (2010). Credibility of the MSP results
271 were checked by three parameters: thermal-induced alteration $|\varepsilon_{\text{alt}}|$ parameter (Fabian &
272 Leonhardt 2010; Monster et al. 2015b); the maximum allowed angle (θ) between the isolated
273 NRM and acquired pTRM; and the intersection parameter (Δb) (Monster et al., 2015b), which
274 tests whether the linear fit regression line intersects the y-axis at the theoretically predicted
275 value of -1 . In order to confirm the obtained and only reliable results, we have set $|\varepsilon_{\text{alt}}| \leq$
276 5%, θ must be less than 10° , and a threshold value of Δb is ± 0.1 . In the MSP-Tool software,
277 the bootstrap statistics were applied to calculate the mean and 95% confidence intervals.

278 Technically, successful MSP experiments were obtained from profiles V and HT, while both
279 HM and HB failed to give reliable estimates as their ε_{alt} parameter exceeded the defined
280 limit. In V, three specimens out of nine were rejected (Fig. 5a) as they have altered during
281 the MSP run. The remaining six met the criteria limit defined above and thus a domain state
282 corrected PI mean of $62.9 \pm 2.6 \mu\text{T}$ was obtained for the vertical profile. The HT gave

283 successful MSP experiment in eight out of ten specimens (Fig 5b), with PI value of $58.6 \mu T$,
284 after DSC procedure. We note here that 95% confidence interval in HT ($+6.5/-6.3 \mu T$) is
285 almost double the value of the confidence interval in V. Obviously, this high scatter is
286 reasoned by the noticeable nonlinearity behavior, in the last field steps (Fig. 5b), between the
287 TRM and applied magnetic field. This non-ideal behavior can be attributed to presence of
288 large ferromagnetic particles, which can reduce the efficiency of linearity law (Selkin et al.,
289 2007). Regarding to HM, only one specimen out of nine passed the alteration limit (Fig. 5c),
290 indicating that most of the middle-zone specimens are susceptible to alteration. In addition,
291 the last data points are not aligned linearly which probably point to the dominance of MD
292 particles. Therefore, no reliable results were obtained from this profile.

293 We have neglected the criteria limits to obtain reliable results without regard to data quality,
294 just to see the impact of the sample position in vertical profile on the PI results. A MSP-DSC
295 value of $67.2 \mu T$ was obtained, this is demonstrated in the supplementary Figure S1. Three
296 accepted data out of nine (33% success rate) was found along HB profile (Fig. 5d) and thus
297 no meaningful estimate could be obtained. As in HM, if we neglect the criteria set parameters
298 (Fig. S2), a value of $60.8 \mu T$ is calculated for HB profile.

299 From the above results, it is obvious that only the vertical and upper horizontal profiles gave
300 reliable results that meet the proposed criteria limits. Therefore, combining all accepted
301 specimens from V and HT enables us to assign Xitle-mean PI (Fig. 6a) at $60.5 (+4.0/-4.1)$
302 μT . Including all the 18 specimens of HM and HB (Fig. 6b) gives, unexpectedly, a mean
303 value of $63.8 (+5.2/-5.8) \mu T$, which considering the uncertainty limits, is indistinguishable
304 from the mean value calculated from only reliable results. This consistency may denote that
305 the current conditions set for accepting the MSP results (ϵ_{alt} and Δb) are ineffective. Such

306 explanation, however, cannot be confirmed in this study, as the actual intensity value during
307 Xitle eruption time (370 AD) is unknown. But, we can deduce that obtaining reliable MSP
308 results for a certain lava unit could be achieved by taking many specimens as possible from
309 different parts within the lava flow. However, these approaches will cost effort and time.

310 **6. Discussion**

311 Three horizontal profiles were sampled from Xitle in three levels from bottom to top in order
312 to discuss horizontal variations of paleointensity. Horizontal characteristics along these
313 profiles were investigated as well through sampling one vertical profile of *ca.* 4.5 m
314 thickness. Rock magnetic experiments were completed to infer the type and size of the
315 ferromagnetic minerals and their thermal stability. It must be mentioned that the number of
316 present samples that have undergone magnetic experiments and MSP run are few and uneven,
317 and thus the relationship between magnetic properties and paleointensity behavior along
318 lava's profile may not be clear. However is not the main focus of this study, as we here try
319 to enhance the PI estimates of Xitle by evaluating previous data and conducting a new MSP
320 experiment. Completing this, we can provide an average value with low confidence limit,
321 through combining present PI results with reliable results obtained from previous studies.
322 Despite the limited number of data, we can give a general overview of the impact of lava
323 position and its physical and magnetic properties on the MSP results. Currently, few studies
324 have addressed the relation between rock magnetic properties and lava flow thickness. Some
325 studies (Coe et al., NATURE, 1995; Rolph, 1997; Hill and Shaw, 2000; V  rard et al., 2012)
326 have discussed this relation on thin lava flows, thickness < 2 m, some others extended these
327 studies to thicker lavas of ~6m long (B  hnel et al., 1997; de Groot et al., 2014), and up to
328 several tens of meters (e.g. Wilson et al., 1968; Audunssen et al., 1992). Findings from these

329 studies indicate an effect of the sampling location on the magnetic properties. Lava thickness
330 reflects different cooling history, the top and bottom parts cooled faster than the central part
331 of the flow. These variations in cooling time govern the size of the ferromagnetic crystals
332 and their oxidations states. The top part of the lava flows may produce smaller ferromagnetic
333 sensus lato particles with lower oxidation state in comparison to those formed in the middle
334 of the flow (see for example Böhnel et al., 1997; de Groot et al., 2014).

335 Böhnel et al. (1997) showed detailed results of Xitle's rock magnetic properties (i.e. Tc,
336 magnetic susceptibility, hysteresis parameters, and coercivities) and PIs along a vertical
337 profile of 6 m thickness. Results indicate that, unlike magnetic properties, the PIs seem to
338 have a systematic behavior with tendency of middle flow samples to give larger PI variations
339 in comparison to the top and bottom samples. Despite the detailed study, they did not find
340 meaningful relation between PI variations and physical and magnetic properties. Our
341 sampled profile spaced only (0.96 km) from Böhnel et al. (1997) profile, implying that they
342 shared almost the same cooling history. The rock-magnetic experiments done in this study
343 showed that Ti-poor T_{Mt} and/or M_t is the dominant ferromagnetic mineral(s) and few
344 hematite is present, showed from *k*-T curves (Fig. 4a). These minerals are located broadly on
345 the PSD size with wide range of MDF (from *ca.* 5 to 56 mT). From top to bottom (V), and
346 along each horizontal profile (HT, HM, HB), we do not find any systematic changes on the
347 magnetic properties although of the, above mentioned, differences in cooling times. Although
348 the magnetic parameters obtained in this study varied (e.g. RP%, MDF and MDT)
349 horizontally, we consider that they do not reflect true magnetic properties changes, as the
350 cooling time and content of the cooled lava on such a small scale (1.5 m) are seemingly
351 constant, at least in comparison to vertical direction. This discrepancy in magnetic data could

352 be attributed to the limited number of rock magnetic experiments carried out. In conclusion,
353 there is not a direct connection found between lava thickness and its 1.5 m lateral extent with
354 the magnetic properties.

355 Now, we discuss the influence of cooling time on the MSP results. Our contributed results
356 show that the topmost horizontal profile and the vertical profile have yielded successful MSP-
357 DSC PI determinations from the 14 selected samples out of 19 available (74% success rate).
358 On the hand, samples from central and bottom parts of the flow do not give satisfactory MSP
359 results, as they have exceeded the limits of criteria set designed following Monster et al.
360 (2015). Apparently, getting differential MSP results and success rate from over a flow section
361 of 4.5 m thickness can be directly related to the cooling time of this flow. It means that top
362 part of the Xitle flow is appropriate for conducting MSP experiments as it has cooled faster
363 than underlying horizontal sections. Moreover, conducting the experiment vertically does
364 give almost the same results given from HT profile (Fig. 5a and b), which could indicate that
365 the effect of cooling time on MSP results may be insignificant or even disappeared if samples
366 are taken vertically. This explanation is based on a few number of samples and more studies
367 are needed to emphasize it. From these notes, we recommend sampling a lava flow
368 horizontally from its top part when the lava flow is not covered by a younger flow, and
369 vertically in order to give a reliable MSP estimates.

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371

372 **7. Conclusions**

373 The rock magnetic properties indicate that the main magnetic minerals are Ti-poor
374 titanomagnetite with small contribution from hematite of PSD carriers accomplish the quality

375 criterion to be able for paleointensity experiments.
376 Successful MSP-DSC experiments were obtained from profiles V and HT, while both HM
377 and HB failed to give reliable estimates. Results obtained in this study from 14 accepted
378 specimens of HT and V profiles give an average mean for Xitle at $60.5 (+4.0/-4.1) \mu T$. This
379 value constrains greatly the dispersion and we consider it should substitute the value for the
380 Xitle from the database used by models in order to do more precise the secular variation
381 curve used for the dating of geologic and archeomagnetic materials of this period.
382 This mean value is consistent with the mean ($62.0 \pm 9.3 \mu T$) calculated from 134 filtered
383 Thellier & microwave old PI data. The whole PI mean value for Xitle, using a combined
384 formula (Higgins and Green, 2011) gives $61.9 \pm 9 \mu T$. This mean value is calculated from
385 high quality data provided from three different PI methods, therefore most likely represents
386 the intensity value for Central Mexico at *ca.* 370 AD.

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388

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561 List of Figures

562

563 **Figure 1.** Distribution of Xitle lava flows I to VI, with the location of the sampling site and
564 the Cuicuilco archeological site. Modified after Delgado et al. (1998).

565 **Figure 2.** Evaluation of previous PI data published for Xitle. The closed and open circles
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579 **Figure 3.** Sampling profile. Lava flow with scale (1m). Three zones are observed with the
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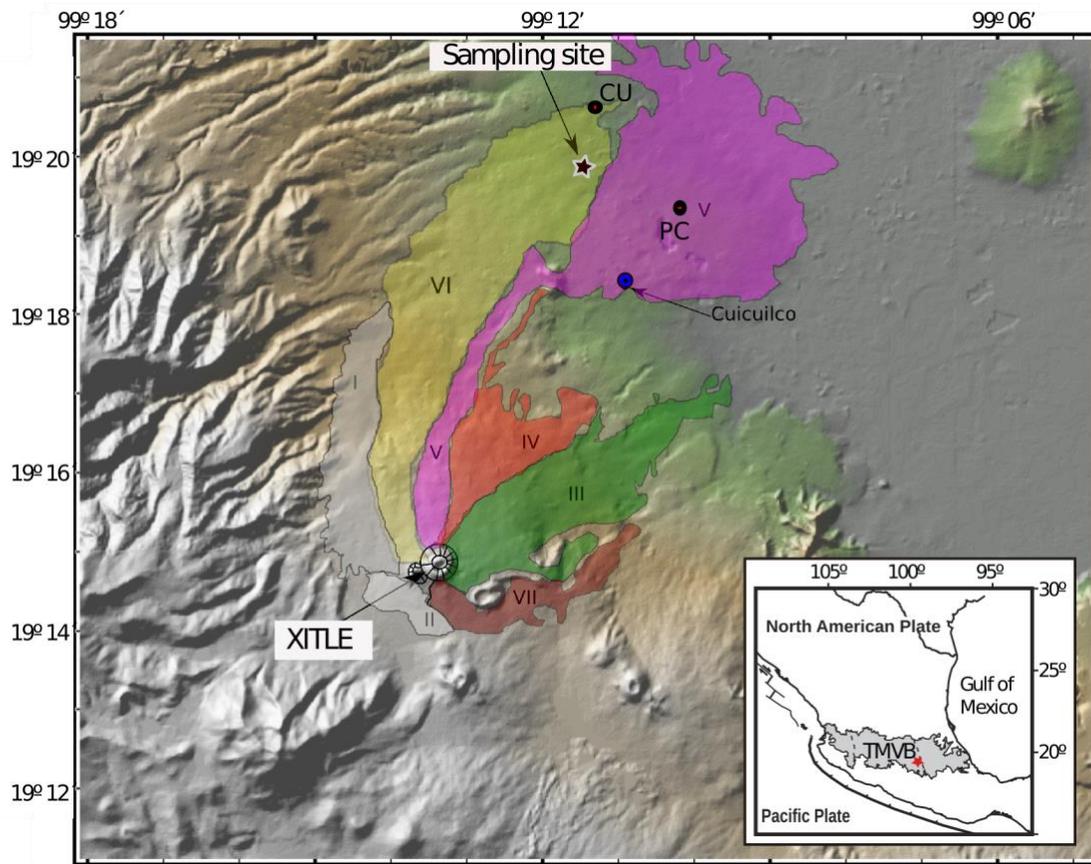
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596 that data were analyzed with MSP-Tool software (Monster et al., 2015b) where bootstrap
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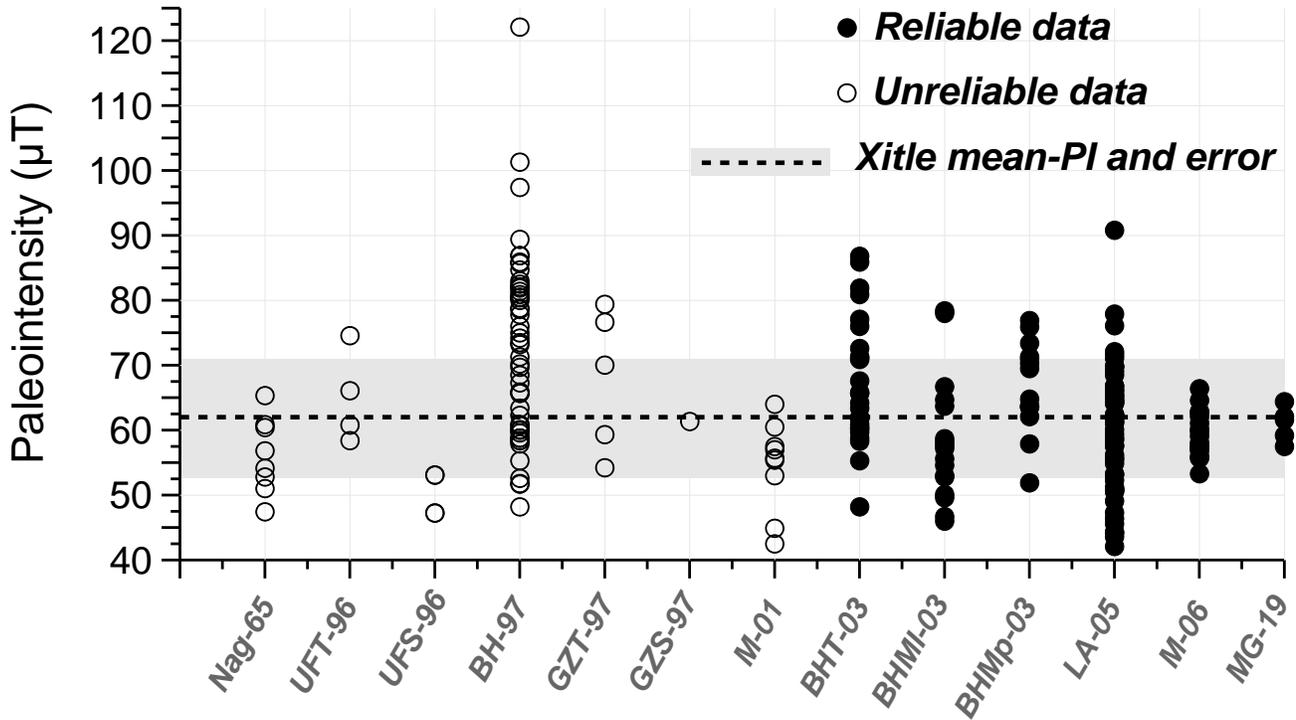
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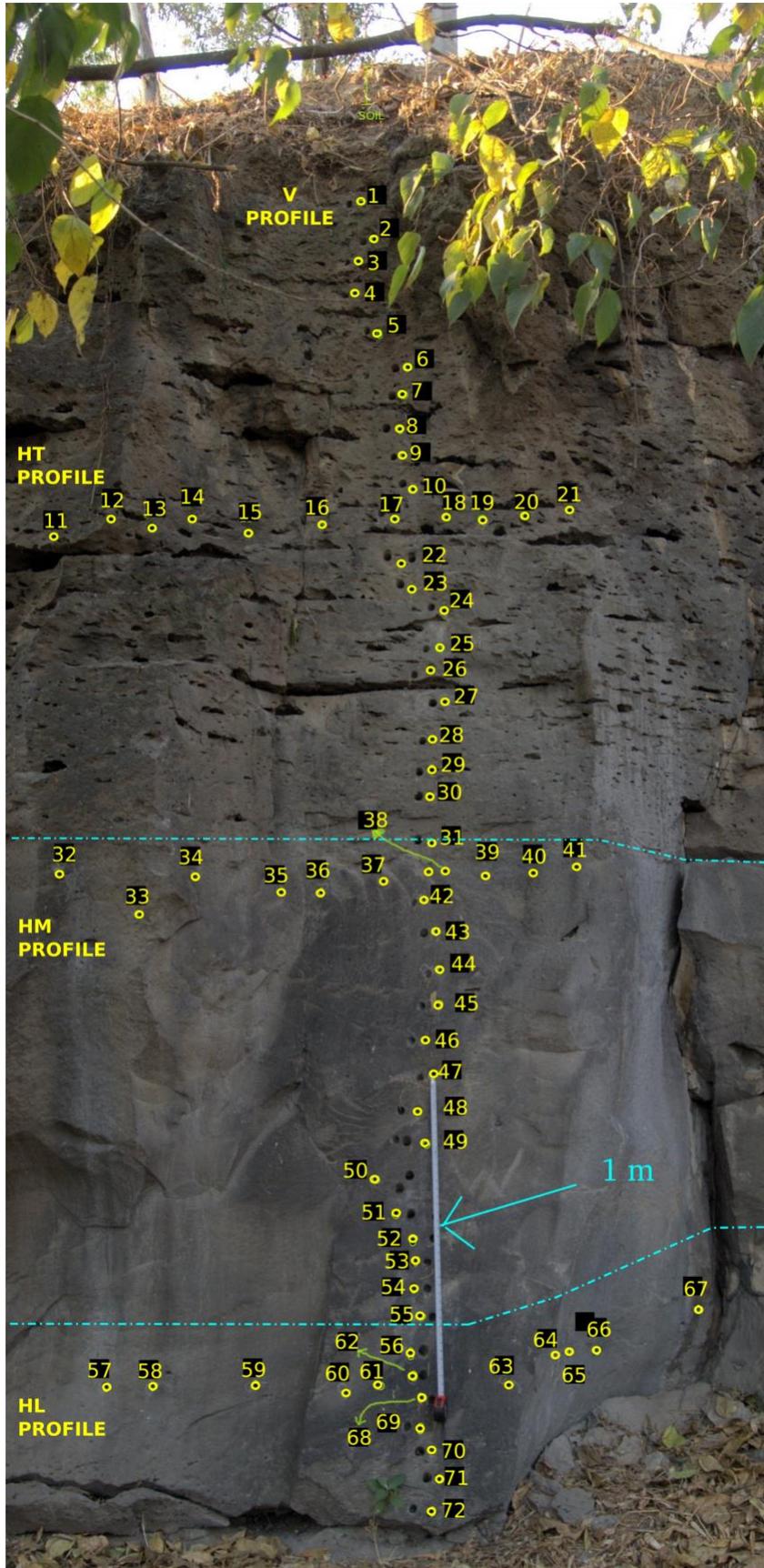
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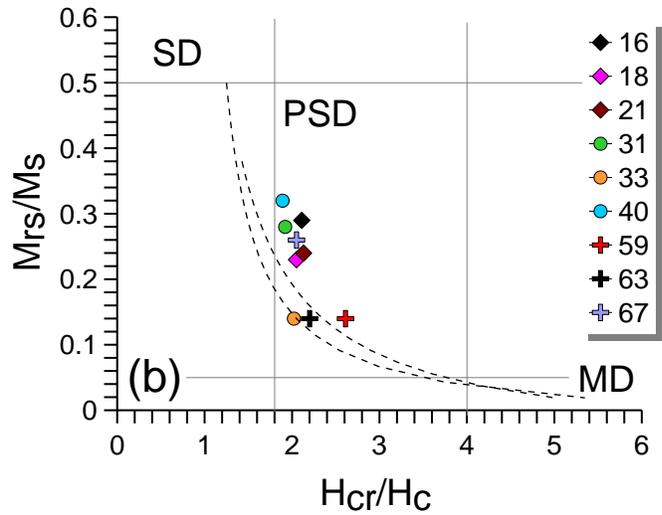
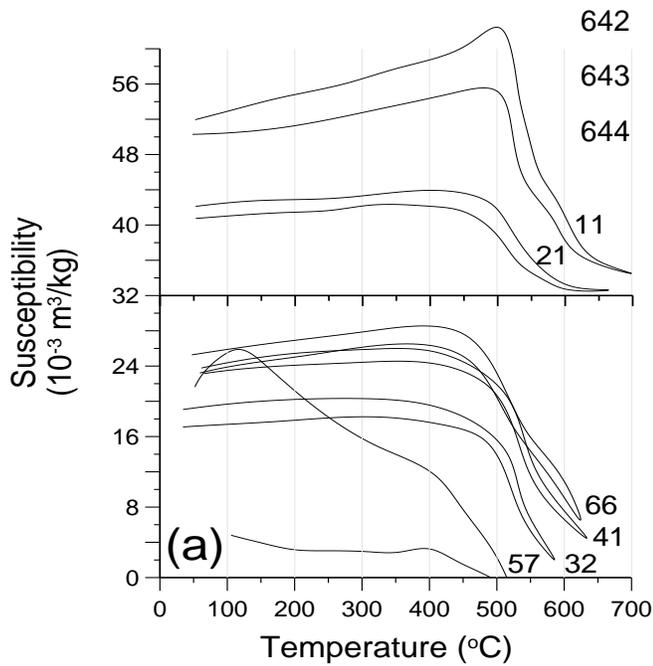
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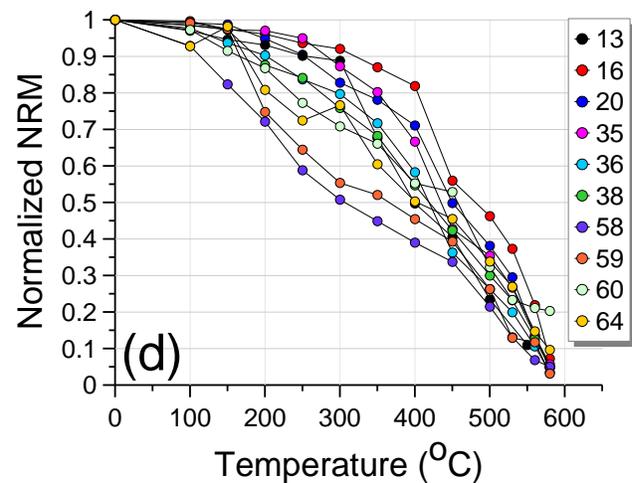
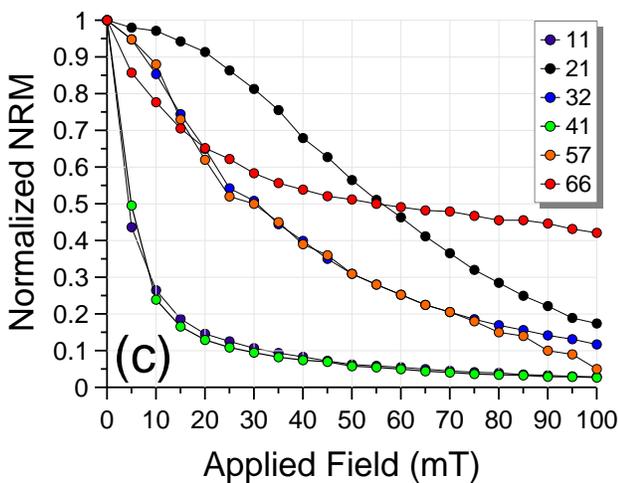
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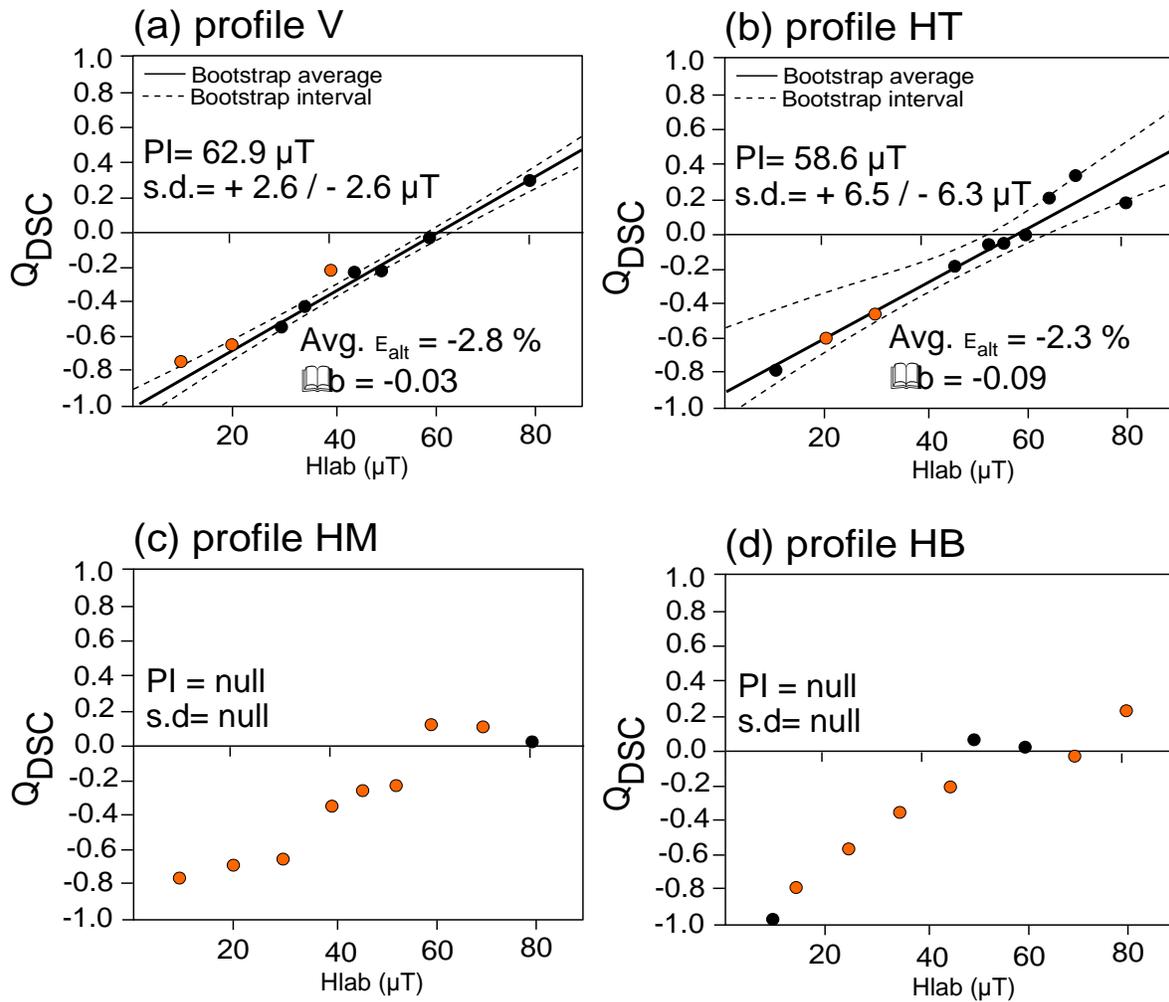


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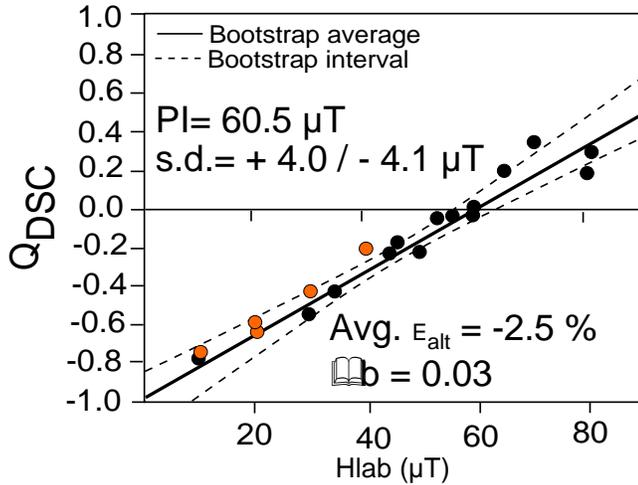
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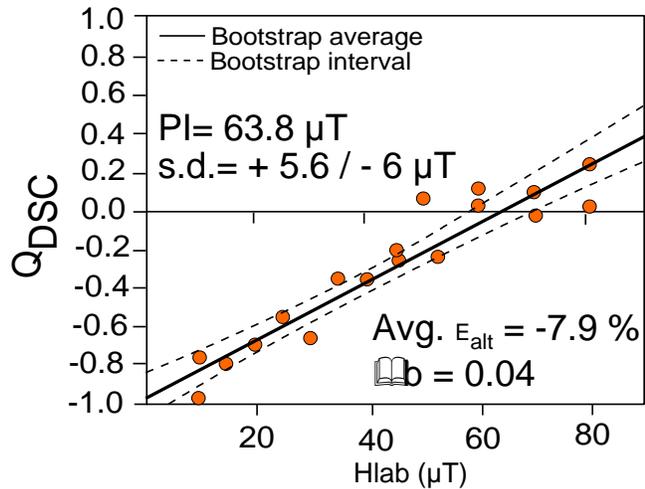
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(a) Reliable experiment:
V+HT



(b) Unreliable experiment:
HB+HM



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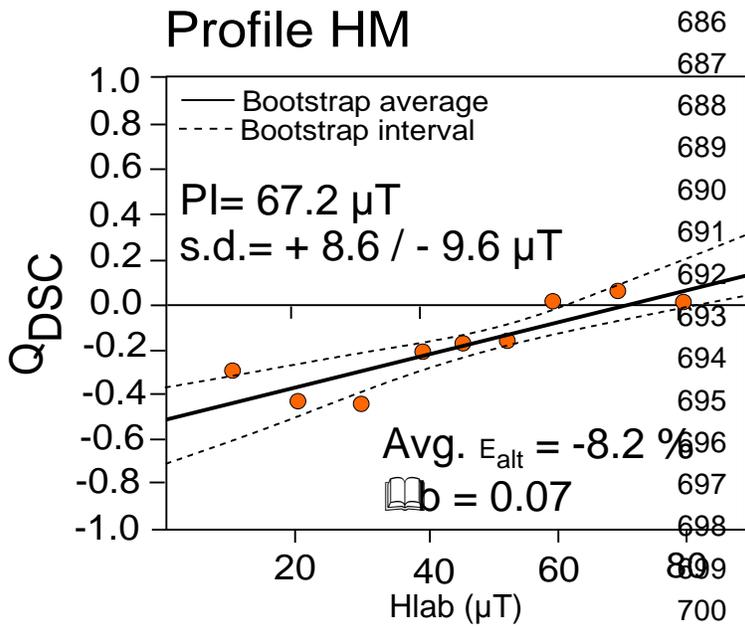
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Supplementary Materials

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684 The supplementary materials consist of two figures.

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702 Figure S1. Multispecimen results obtained from central profile (HM), after domain state
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706 the selection criteria defined in the present study, as we just need to compare the unreliable
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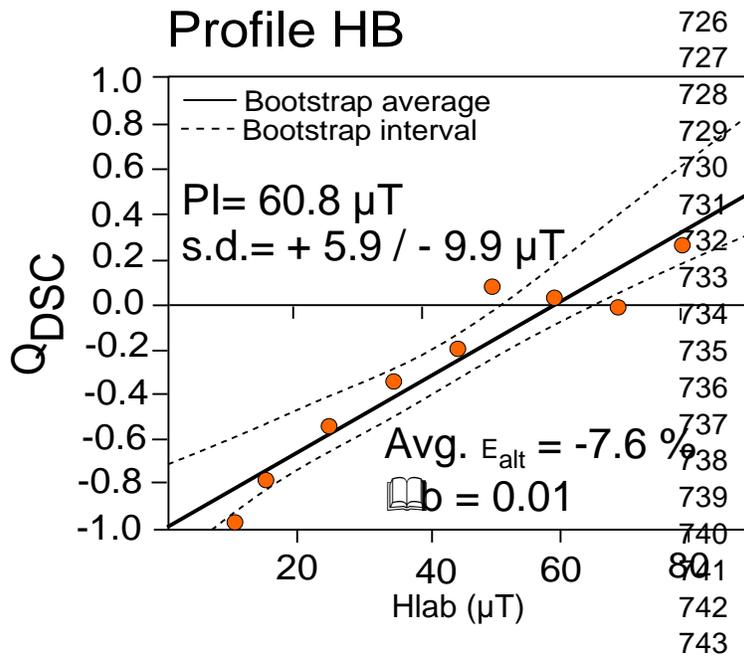
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744 Figure S2. Multispecimen results obtained from bottom profile (HB). For details see Fig.
745 S2 caption.

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