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Sedimentary evolution and ecosystem change in Ahémé lake, south-west Benin

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Abstract. Tropical moist ecosystems, such as Ahémé lake, south-west Benin, are increasingly marked by water degradation, linked with the activities of increasing riparian populations. The objective of this study is to analyze sedimentary dynamics and its influence on the changing ecosystem of Ahémé lake from 1961–2010. Data used to carry out the study are records of precipitation, flows, turbidity, suspended sediment, mineral elements and bathymetry. Grain size data from the sieving of sediment samples were used to interpret suspended solids distribution in the lake. Linear correlation coefficients were used to assess the degree of dependence between rainfall and runoff inputs to the lake. Lake depth measurements in some areas of the lake serve to determine the rate of infilling. The sorting index was used to highlight the distribution and origin of sediments in the lake. The results show a degradation of the lake Ahémé ecosystem characterized by infilling of its bed, a high correlation ($r = 0.90$) between rainfall and runoff, seasonal change in physicochemical parameters (total suspended sediment decrease by $-91\%$) and decrease in fish production by $135.8\ t\ yr^{-1}$. The highest mean suspended sediment concentrations in lake inputs occur during high water periods ($123\ mg\ L^{-1}$) compared to low water periods ($11.2\ mg\ L^{-1}$).

1 Introduction

Ahémé lake, in Benin western Africa, and its biodiversity attract enormous interest from people, local residents and scientists in particular (Amoussou et al., 2016). The riparian population interest is in the supply of fish resources but the ecosystem of Ahémé lake is in constant degradation. The lake dynamics and its water resources are linked to climatic and hydrological variability (Amoussou et al., 2007). Analysis of the dynamics of rivers and water bodies (Vissin, 1998) is necessary to meet the main objectives of the Global Energy and Water Experiment (GEWEX) and the Tropical Atmosphere and Hydrologic Cycle (CATCH) projects.

Eutrophication of Ahémé lake (Oyédé, 1991; Amoussou, 2004) leads to siltation, excessive concentration of chemical elements, resulting in sometimes a significant presence of algae or aquatic plants, and also absence of certain plant species along the lake shores inducing hydraulic erosion.

These combined pressure factors could lead to destruction of the ecological habitats of fish species and consequently disrupt socio-economic activities. Commercial fishing is declining as a result of lower fisheries yields. Thus, the balance between natural resources and human population demands or needs is being compromised. This study aims to analyze the evolution of sedimentary and environmental parameters from 1961 to 2010 and their impacts on the ecosystem of Ahémé lake. This work brings together hydrometric, sedimentary and water quality information to help understand ecosystem changes and its impacts in Ahémé lake.
2 Data and methods

2.1 Study site

Ahémé lake is located between 6°20’ and 6°40’ N and 1°55’ and 2° E (Fig. 1a).

Ahémé lake is located in a depression between the Allada and Comè plateaux (Fig. 1b). The Couffo River flows into Ahémé lake from the north. Ahémé lake exchanges water in the south with the Mono and Sazoué rivers which are connected to a coastal lagoon and the Atlantic Ocean through the 24 km long Aho Channel. During the wet season freshwater from Ahémé lake flows southwards in the channel to the lagoon and Atlantic Ocean. However, during the dry season the flow in the channel reverses, resulting in increased salinity in the southern part of the lake. The width of the lake is ~ 3.4 km at the latitude of Guézin. The surface area of Ahémé lake is between 70 and 100 km² in the dry and rainy seasons, respectively (Le Barbé et al., 1993).

Ahémé lake is influenced by a rather dynamic lagoon system that favors variations in pH, temperature and salinity due to its opening on the Atlantic Ocean. The contribution of Mono River to the lake is most significant during floods or periods of high water levels (Pliya, 1980; Oyédé, 1983; Amoussou et al., 2007).

2.2 Data

Rainfall data from stations at Athiémé, Grand Popo, Bopa, Allada and Ouidah (see Fig. 1a) over the period 1961–2010 were extracted from the database of METEO BENIN. Flow data records of the Couffo River at Lanta and Mono River at Athiémé over the period 1961–2005 (data are not available for 2006–2010) were collected by the Hydrology Department of the Directorate General for Water. Information on sediment dynamics in Ahémé lake was available as: (1) sediment depth data collected only in 1991 and 1999 were extracted from Oyédé et al. (2007); (2) annual suspended solids concentrations measured in 1999 and 2007 at Guézin (Roche International, 1999; Amoussou et al., 2007). These were complemented by seasonal (rainy and dry) bathymetry measurements in 2003 and 2006 in Ahémé lake. Physicochemical measurements were made at localities on the east and west shores of the lake in: (1) October 2000 – pH only; and (2) September 2002 – salinity and pH measured using a WTW 340i handheld pH/conductivity meter. Sediment samples collected during hand dredging were analyzed by sieving using the AFNOR (French Standardization Association) method (AFNOR, 1996). Sediment weighing was carried out on a Shimadzu BX3200D (dual range: 3200/600 g and resolution: 0.1/0.01 g).

2.3 Methods

Linear correlation coefficient were calculated between mean monthly rainfall at Athiémé, Grand Popo, Bopa, Allada and
Ouidah station and mean monthly flow at stations Athiémé and Lanta for the period 1961–2005 (see Figs. 2 and 3). Bravais-Pearson’s correlation coefficient is calculated with 0.05 significance level.

Assessment of the sedimentary evolution of Ahémé lake was conducted to determine whether sediment depositions or erosion was occurring or whether there was sedimentary balance. The lake cross-sectional area is calculated from the depth and width of the wetted section. Thus, the decrease or increase in section area is related to depth, because from one year to the next in the same season, the width varies very little.

The bathymetry is measured by boat over 2.5 km at Bopa Kpindi, Bopa Centre and Ségbobouè (Fig. 1b), from the east to west bank. On this cross section, the measurement depths of Ahémé lake is doing each 2 m distance with graduated wooden ruler on dry and rainy seasons.

These measurements referred to a standard water level (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m) at the gauging station under Guézin bridge (Fig. 1b) and formed the basis for estimating the sediment depth data (0.5 m).

\[
S_0 = \frac{q_3}{q_1},
\]

where \(q_1\) and \(q_3\) are the first and third quartile, respectively, of the grain size distribution.

- If \(S_0 = 1\) or close to 1: the sediment is homogeneous (well sorted), corresponding to a steep gradient of the grain size curve (Ben Amor et al., 2003; Marc and Emblanch, 2005);

- If \(S_0 < 1\) or \(> 1\): the sediment is poorly sorted out, corresponding to a low gradient of the grain size curve (Ben Amor et al., 2003; Marc and Emblanch, 2005).

Figure 2. Mean monthly rainfall at the stations around Ahémé lake over the period 1961–2010.

Figure 3. Mean monthly flow of the Mono River at Athiémé and Couffo River at Lanta linking Ahémé lake over the period 1961–2005.

3 Results and discussion

3.1 Rainfall and river flow variation

Figures 2 and 3 show the mean monthly rainfall and flows of rivers feeding the Ahémé ecosystem. The annual rainfall regime is bimodal, characterized by an important supply of water over May–July (rainy season) and September–October (small rainy season).

This seasonal climate variability results in the transport of dried sediment from the banks into Ahémé lake. Sediment and water inflow from the Couffo River to Ahémé lake is less important than that of the Mono River which has a higher flow (annual mean flow is 5.1 m³ s⁻¹ at Lanta and 114.4 m³ s⁻¹ at Athiémé over 1961–2005).

Correspondence of flows with the rainy season confirms the high rainfall vs. flow Bravais-Pearson’s correlation coefficient \((r = 0.90)\) in the study area.

Moreover, the lake receives a large volume of water during September–October when the rivers entering the lake reach their maximum annual flow and also contain their highest suspended solids concentration as reported by Amoussou (2004). The Couffo river has much greater seasonal variability in flow compared to the Mono river especially since construction of the dam on the Mono river at Nangbéto (7°25'25.40"N; 1°26'5.82"E) in September 1987. The difference in the hydrological flow regimes in the Couffo and Mono rivers and the effect of the Nangbéto dam on dry season flow are evident in Fig. 3. Base flow in December to April at Athiémé was 4.22 m³ s⁻¹ before the dam construction and increased to 57.26 m³ s⁻¹ after the dam construction. The dam has had a major impact on water and sediment flows from the Mono river to the Ahémé lake ecosystem, including increased bank erosion downstream of the dam (Oyédé, 1991; Amoussou, 2010).

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from the sea. Soluble materials from fine particles, indicating the transport of particles in suspension. These results are consistent with those of Cerdan et al. (2002) and Amoussou (2010), who reported that, because of the absence of vegetation, in the rainy season sediment particles are mobilized and deposited on the lake bed, modifying the lithofacies of the bottom. Large sediment particles are deposited on the river banks whilst finer sediment particles are transported by the river into Ahémé lake.

Values of the Sorting Index ($S_0$) were $> 1$, showing that deposited sediments are poorly sorted in both Ahémé lake ($S_0 = 2.15$) and in the lagoon ($S_0 = 1.71$). This can be explained by the effect of a range of processes, including low roughness of the floodplain, flocculation due to saline conditions at high tide, solid inputs generated by anthropogenic activities and the transport and deposition of the majority of sediment as bed load mobilised during the rainy season. These results are consistent with those of Cerdan et al. (2002) and Amoussou (2010), who reported that, because of the absence of vegetation, in the rainy season sediment particles are mobilized and deposited on the lake bed, modifying the lithofacies of the bottom. Large sediment particles are deposited on the river banks whilst finer sediment particles are transported by the river into Ahémé lake.

### 3.4 Evolution of physicochemical parameters

Salinity measurements at the entrance to the lake (Guézin first bridge) in the north (Bopa Agonsa) and on both banks of Ahémé lake in September 2002 (Fig. 5) show high salinity even during the rainy season. This is caused by the quasi-permanent opening of the Mono River mouth on the coastal lagoon. It could also be attributed to a decrease of rainfall around the lake in the years 2000 ($-26$ to $-29\%$), 2001 ($-8$ to $-13\%$) and 2002 ($-10$ to $-15\%$), compared with the annual mean rainfall for the period 1961-2010 (968.42 mm), providing less rainfall for dilution of salinity. Salinity values measured in the lake in September 2002 (at the beginning of the floods of the Mono and Couffo rivers, where salinity is almost zero) are much higher than those recorded in other parts of the drainage basin (Oyédé, 1981, 1983, 1991) and other rivers in Benin (Amoussou, 2003; Amoussou et al., 2007).

Increasing salinity can result in enhanced flocculation of terrigenous sediment in the lake and subsequent deposition within the lake, contributing to sediment infilling.

Analysis of the pH data (Fig. 6) allows assessment of the suitability of the lake water quality for aquatic ecology and fish species. The mean pH at the localities ranges from 7.2 to 8.2 in October 2000, indicating a basic environment dur-

### Table 1. Sedimentation evolution in some parts of the Ahémé lake from 1991 to 1999.

<table>
<thead>
<tr>
<th>Site</th>
<th>Sectional areas of sediment (m$^2$)</th>
<th>Difference (1991–1999)</th>
<th>Variation rate (%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Bopa Kpindji</td>
<td>2405–2499</td>
<td>+94 (3.9)</td>
<td></td>
</tr>
<tr>
<td>Bopa Centre</td>
<td>5164–4919</td>
<td>−245 (−4.7)</td>
<td></td>
</tr>
<tr>
<td>Ségbohouè</td>
<td>2899–2744</td>
<td>−155 (−5.3)</td>
<td></td>
</tr>
</tbody>
</table>

Source: Oyédé et al. (2007)

### 3.2 Morphodynamics of Ahémé lake

Data from the years 1991 and 1999 (Oyédé et al., 2007) revealed declining water depths in some areas of Ahémé lake. Comparison of the depths between 1991 and 1999 in a few areas of the lake, in the north (Bopa Kpindji and Bopa Centre) and the south (Ségbohoué), revealed infilling (Table 1) from the east bank to the west bank. In the north of the lake, there was erosion at Bopa Centre from 1991 to 1999, together with deposition at Bopa Kpindji, especially on the eastern bank.

Infilling of Ahémé lake, as indicated by some of the data presented here, is one of the factors of the lake impoverishment contributing to fish species migration and ecosystem degradation.

Suspended solid inflows to the lake are more important in the rainy season than in the dry season, as demonstrated by measurements carried out by Roche International (1999) and Amoussou et al. (2007), who showed that at Guézin bridge, the daily suspended solids concentration measured during the rainy season varied from 81 to 165 mg L$^{-1}$ (mean $\sim 123$ mg L$^{-1}$) while in the dry season they range from 3.5 to 19 mg L$^{-1}$ (mean $\sim 11.2$ mg L$^{-1}$).

### 3.3 Granulometric analysis

Figure 4 shows the parabolic curves of granulometric variation for sediment samples from the coastal lagoon and Ahémé lake, based on the sediment weight data transformed into a cumulative percentage.

The proportion of sands ($> 0.5$ mm diameter) is greater (80%) in the coastal lagoon than in Ahémé lake (40%). The particle size range is narrower in Ahémé lake and is dominated by fine particles, indicating the transport of particles in suspension. These results are consistent with those obtained by Yalin and Karahan (1979) on the secondary tributary of the Loire river and by Fournier (2004) on the Durance river, both in France, and by Degoulet (2006) on embankment dams. The textural heterogeneity of the sediments is also due to their origin: either from the crystalline basement geology or sedimentary rocks affected by erosion, or from the sea.

![Figure 4. Sediment granulometric curves in the coastal lagoon and Ahémé lake (Source: Amoussou et al., 2016).](image-url)
Figure 5. Spatial variation in salinity from South to North on the eastern and western bank of Ahémé lake in September 2002 (Source: Amoussou, 2003).

Figure 6. Spatial variation of pH measured in Ahémé lake in October 2000 (Source: PAZH, 2002).

Figure 7. Annual fish production in Ahémé lake from 1987 to 2000 (Source: Fish farming Directorate, 2002).

4 Conclusion

Ahémé lake is characterized by a bimodal rainfall regime (two dry seasons and two alternating rainy seasons) and the unimodal flow regime of the Mono (with increased base flow as the result of dam construction) and Couffo Rivers. The high flows coincide with the heavy rains occurring in the lake’s catchment. The strong flow/rain relationship reflects the dependence of the flow on rainfall and influences the variation of water physicochemical parameters. Infilling of the lake is confirmed by the continuing decrease in lake depth by sediment with a narrower range of particle size in the lake than in the coastal lagoon.

Data availability. The data are not publicly accessible:

– Climatology data are been propriety of Agence Météo Benin,

– Fish data are been propriety of Fish direction.

Competing interests. The authors declare that they have no conflict of interest.
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